What Did the 1700 AD Cascadia Earthquake Look Like? Correlating Deformation and Tsunami Inundation Modeling with Paleoseismic Proxies

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Goals/Focuses of this work:

- 1. Utilize paleoseismic proxies to study the 1700 Cascadia event
- 2. Devise a method for testing synthetic rupture models using regional paleotsunami data
 - Test the sensitivity of inundation models to tide levels
- 3. Constrain potential slip patterns of the 1700 CE event... kinda.



Why do we care? – implications for the future



- 1. Understanding the earthquake cycle of the past helps our understanding for the future
- 2. Are full margin ruptures possible or likely?

Why do we care? – implications for the future



- 1. Understanding the earthquake cycle of the past helps our understanding for the future
- 2. Are full margin ruptures possible or likely?
- Model A produces wider regions of seismic hazard potentials than Model B

How can we test likelihood of one model over the other?

quence?

What is a paleoseismic proxy?

- Japanese orphan tsunami (26th January 1700 CE)
- Coastal subsidence records (100-400 yrs BP)
- Deep sea turbidites (260 yrs BP +
 - 120 yrs)
- Ghost forest t
- Coastal tsunar

A paleoseismic proxy is a **geological clue** that can be attributed to a past earthquake. It is the **effect** of a rupture recorded as **indirect evidence** of the event







• What did the 1700 event slip

distribution look like?

Paleoseismic proxies:

- Japanese orphan tsunami (26th January 1700 CE)
- Coastal subsidence records (100-400 yrs BP)
- Deep sea turbidites (260 yrs BP +-120 yrs)
- Ghost forest tree rings (1699 CE)
- Coastal tsunami deposits

Summary of tsunami heights, 1700 and 1960





Things we don't to know:

- What did the 1700 event slip distribution look like?
- Single event or sequence?



Methods

Stochastic slip rupture modeling

- Computationally fast kinematic models based on general statistical parameters
- 37,500 ruptures of Mw 7.8 9.5
- Varying rupture area (full and partial margin ruptures permitted)
- Stochastic ruptures are tested based on 3 proxies:
 - 1. Does the rupture produce coastal subsidence that matches the record?
 - 2. Does the rupture produce noticeable tsunami heights at Japan sites?
 - 3. Does the rupture produce inundation that matches coastal deposit records?



LeSelle et al., 2022







Paleosubsidence

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Example stochastic rupture models



Methods

Inundation modeling

- Fine resolution inundation modeling using Geoclaw (https://www.clawpack.org/geoclaw.html)
- Two sites of focus: Salmon River and Alsea Bay, OR
 - Pro: high resolution core data (tsunami deposit thickness estimates)
 - Con: close to one another (70km)
- Include different tide stages to test sensitivity of inundation potentials
 - Highest/Lowest tides (+/- 2m), 50th percentile tides (+/- 0.8m), and zero tide
 - Projected tide level at the time of the January 26^{th} event is ~ -0.8m!



Courtesy of Andrew Meigs, OSU

Comparing coseismic subsidence

- Qualification for matching subsidence:
 - 1. Match coastal subsidence sites located < 50km away from rupture area with RMS < 0.4m.
- Since we are relaxing the assumption of full margin rupture, a single partial rupture does not need to produce all necessary subsidence!
- 1,635/37,500 match local coseismic subsidence to RMS<0.4m





Melgar, 2021

Results

Comparing Japan tsunami

- Qualification for matching tsunami:
 - Produce tsunami that matches records in Japan
 - Or does not produce tsunami > 30 cm 2.
- Since we are allowing partial ruptures to occur, a rupture does not need to produce a noticeable tsunami in Japan to be considered a potential
- 93/1,635 events fit subsidence & tsunami
 - > Mw 8.6
- 529/1,635 events fit subsidence but tsunami < 30cm

For the next phase we subset ruptures down to only those who produce subsidence at the two sites of interest... 230!



Comparing inundation







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Results

Tidal influence in inundation modeling



Tidal influence in inundation modeling





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Comparing Inundation

*technically with such close sites, we cannot rule in or rule out ruptures. But we can provide a framework for doing so later...



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Results

Comparing inundation on January 26th, 1700, low tide



Key points

- Tidal stage plays an influential role in inundation
 - Alsea Bay much more than Salmon River
 - Highest high tide almost always inundates completely
- 7/37,500 ruptures fit all three proxies for the January 26th tide level
 - This includes both partial and full margin ruptures!
 - Mw > 9.0 needed to fit both sites fully
- What influence do other model conditions have on inundation potentials?
 - Paleotopography reconstruction
 - Bed roughness
 - Sea level rise
- More high resolution survey sites along the PNW needed to best constrain potential slip patterns!



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Thank you!

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Cluster inundation modeling

- Method from Williamson et al. (2021)
- 1/3" resolution in undation modeling for a single event takes ~ 1 days
 - Running models for 200 ruptures with cluster method takes ~1 week compared to months!
- Run models at 3" and modify to create pseudo-fine model at 1/3" resolution



Methods

Cascadia tidal stages

 Tsunami arrived at or shortly following low tide in the evening of 26th January 1700 (Witter et al., 2012)





using **5 tide stages** based on tides during 1 year record

Stochastic slip rupture modeling

- Slip on a fault is a random field (Mai & Beroza, 2002)
 - von Karman autocorrelation function best describes ruptures
- 3 variables defined by simple statistical parameters
 - Correlation lengths and Hurst exponent (Melgar & Hayes, 2019)
 - Fault dimensions (scaling law from Blaser et al. 2010)



(Melgar & Hayes, 2019)

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Discussion topics

- 1. What makes a "good" survey site?
- 2. Homogeneous vs heterogeneous surface roughness?
- 3. Landscape reconstruction for paleo topography?
- 4. What can we sediment transport modeling tell us about 1700 tsunami deposits?



1928 T-Sheet for Alsea Bay





Paleoseismic observations in Cascadia



Stochastic kinematic ruptures: The slip patterns

- Analysis of historical earthquakes shows a
 VonKarman correlation function describes slip well
- We have measured correlation lengths on the USGS database of large earthquakes (Hayes, 2017). They depend on fault dimension



Melgar & Hayes (2019)

-128° -126° -124° Melgar, 2021 50° 48° 46° 44° 42° M9.07 40° 30 40 Slip(m)

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Stochastic kinematic ruptures: The background mean

The correlation lengths (and the Hurst exponent) control the variability but we also have to define the "mean" of the slip pattern

We can make several "geophysically informed" choices



Stochastic kinematic ruptures: Fault dimensions

- We also know there is variability in the rupture dimensions
- Earthquakes of the same magnitude have varying lengths/widths. Some times they are long, some times they are short
- We use a log-normal probabilistic scaling law to replicate natural variability (e.g. Blaser et al., 2010) Rupture length (km)
- We don't always select the **entire** -1/m megathrust when making a rupture



Sequences with more than one earthquake



Sequences with more than one earthquake





Goldfinger et al., 2017