

SUBMARINE LANDSLIDES & SEISMOTURBIDITES IN THE CASCADIA SUBDUCTION ZONE

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USGS RESEARCH PLAN

Key scientific questions:

To what extent and how frequently do potentially tsunamigenic upper plate structures rupture with the megathrust?

- Evaluating seismic evidence for slip to the trench with detailed AUV/ROV surveys
- Examining Quaternary deformation in outer wedge
- Characterizing recent deformation on shoreline crossing faults in southern Cascadia

How do along strike variations in the morphology and structure of the overriding plate relate to possible segmentation of the megathrust?

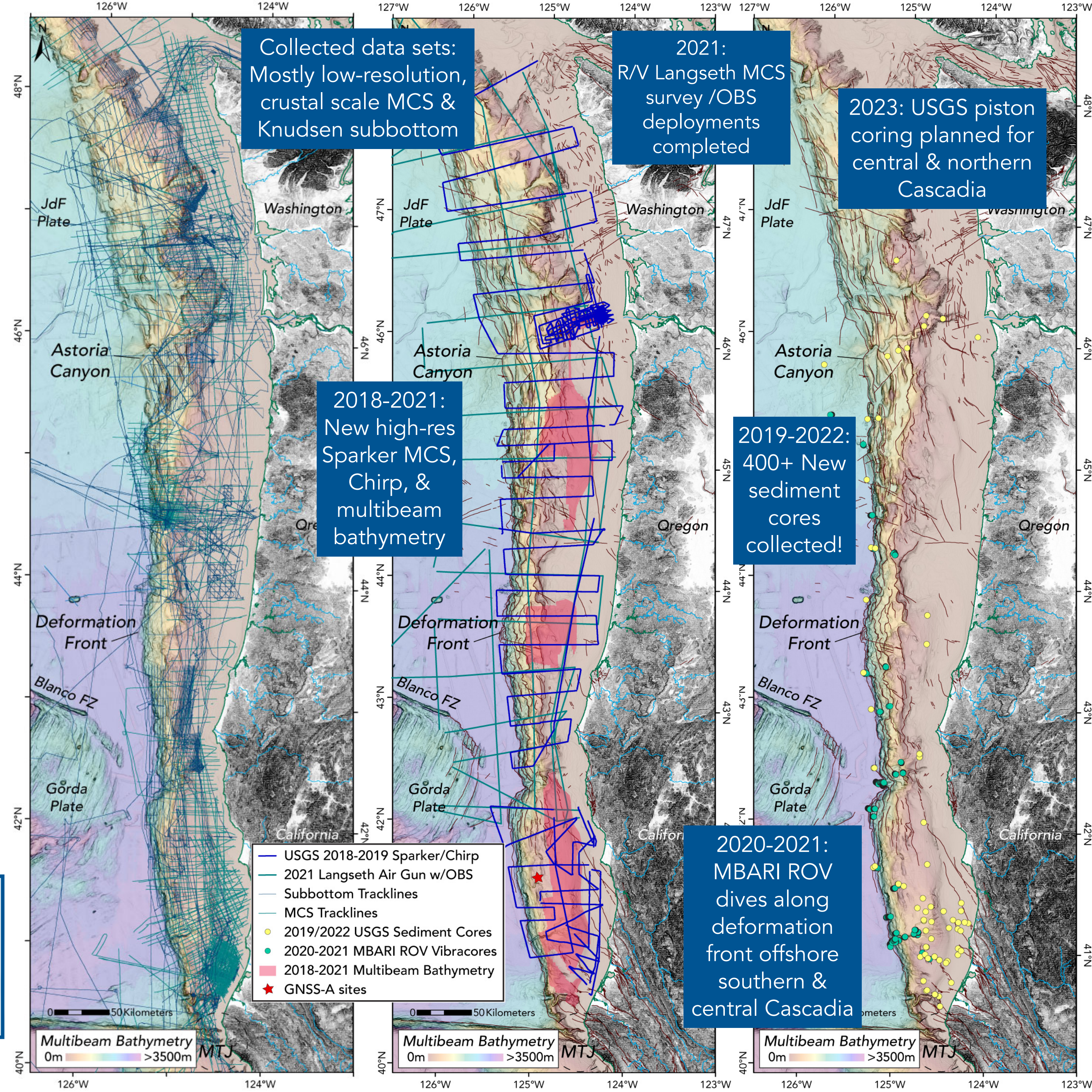
- Watt and Brothers (2021): Systematic characterization of morphotectonic variability...

What is the role of fluids in subduction zone processes?

- Integrating seep mapping with regional structural characterization

How is sediment delivered and redistributed across the continental shelf and slope?

How does earthquake shaking translate to slope failure?

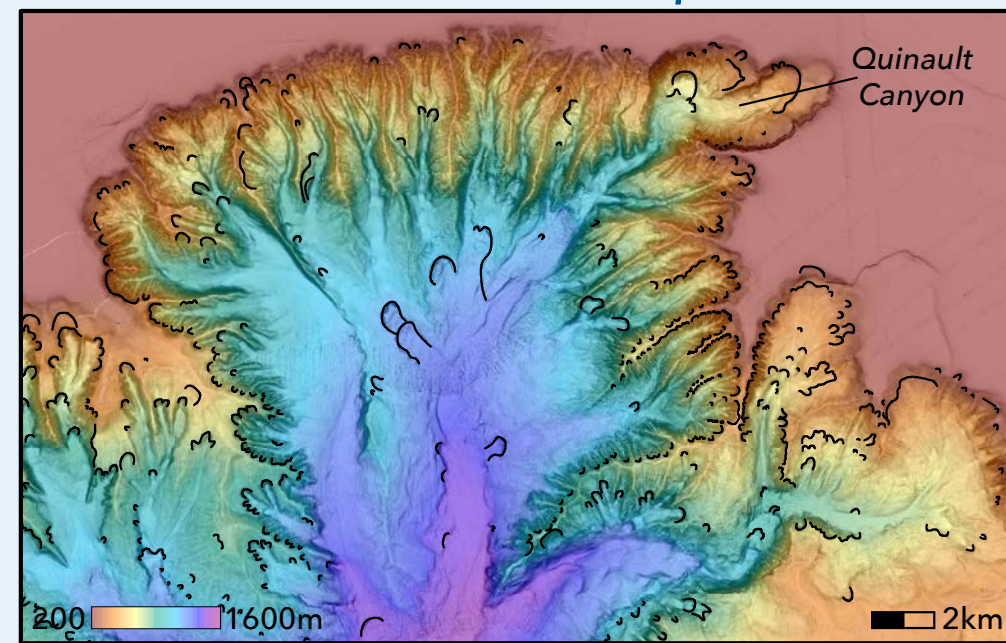


TURBIDITY CURRENTS: SOURCE AREAS & DEPOSITS

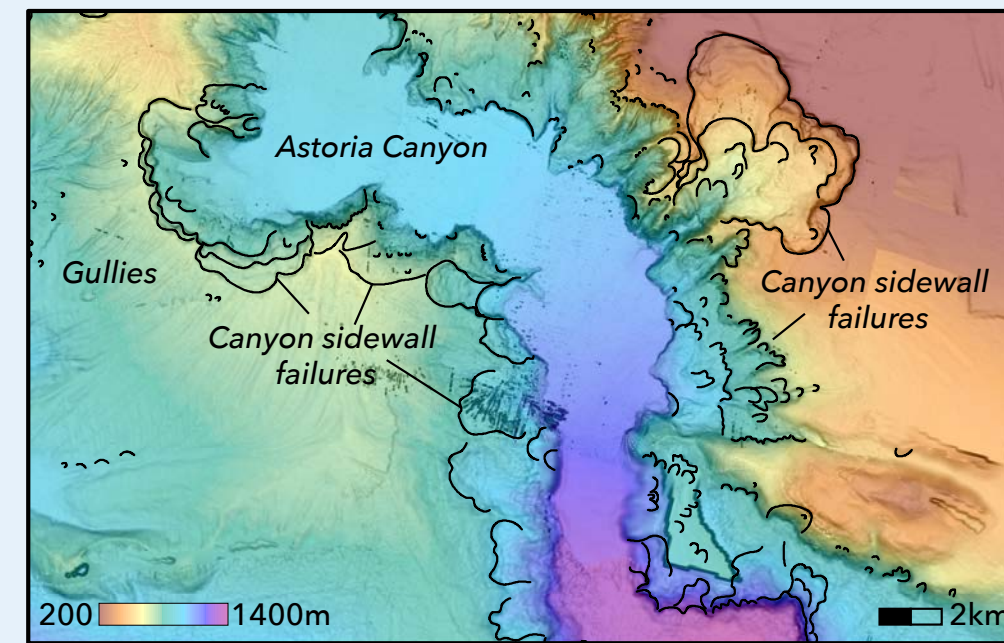
When earthquake shaking occurs, sediment is remobilized along different parts of the slope:

Within submarine canyon systems

Canyon heads with fluvial inputs

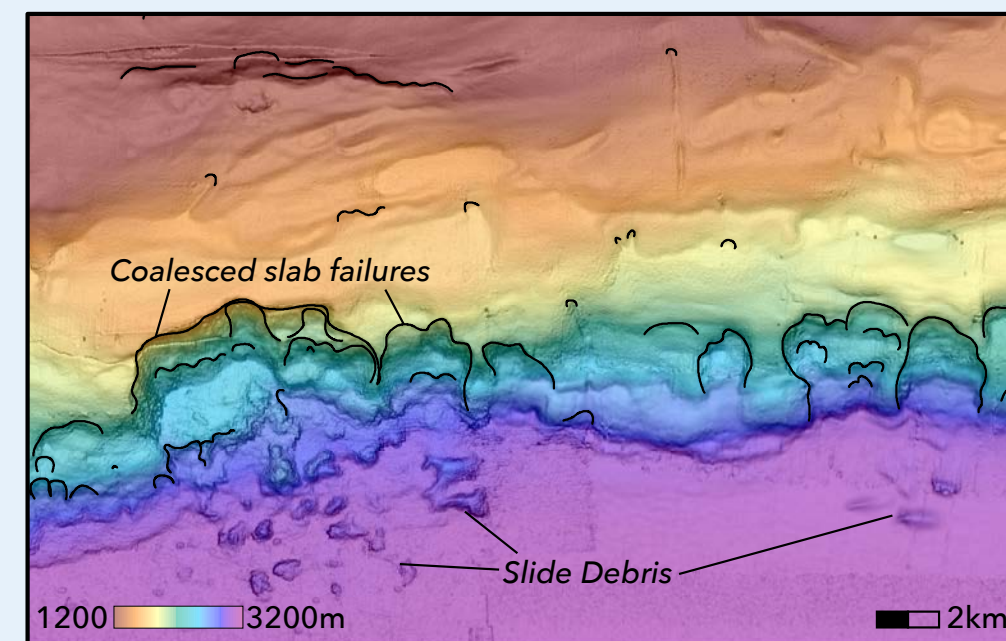


Canyon/channel levees & sidewalls

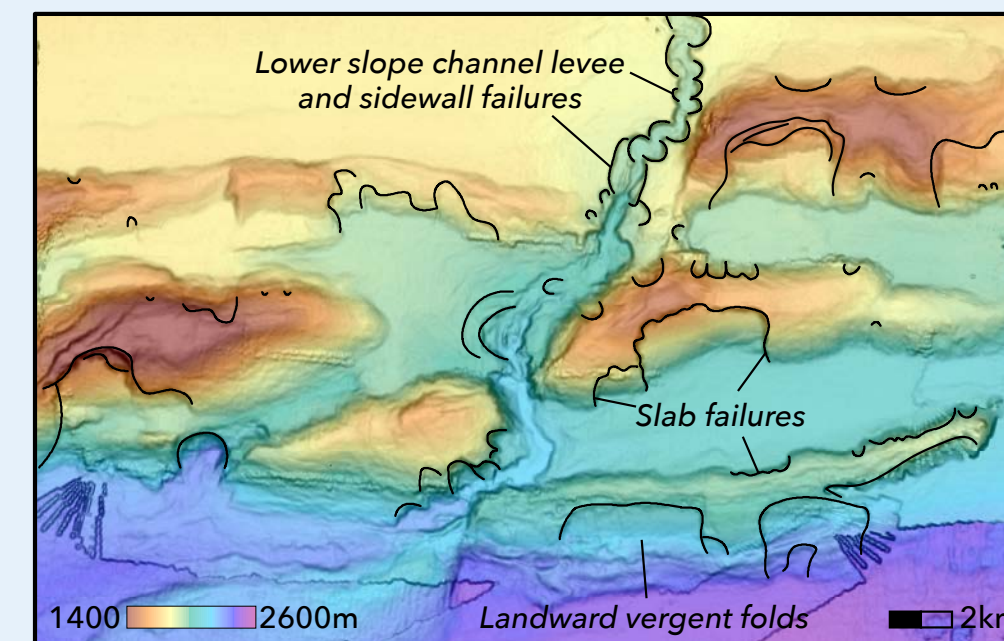


Open slope failures

Steep seaward outer wedge



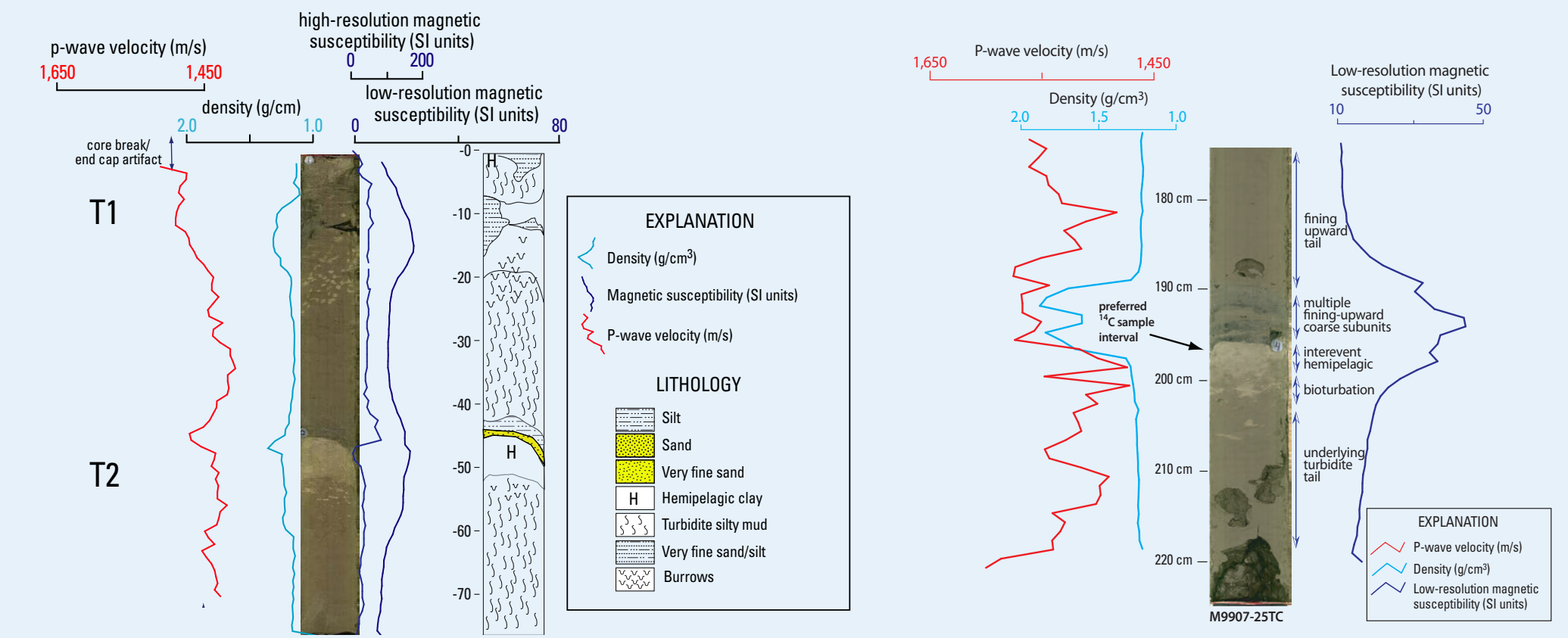
Flanks of anticlinal ridges



From Hill et al., 2022

Sandy abyssal turbidites:

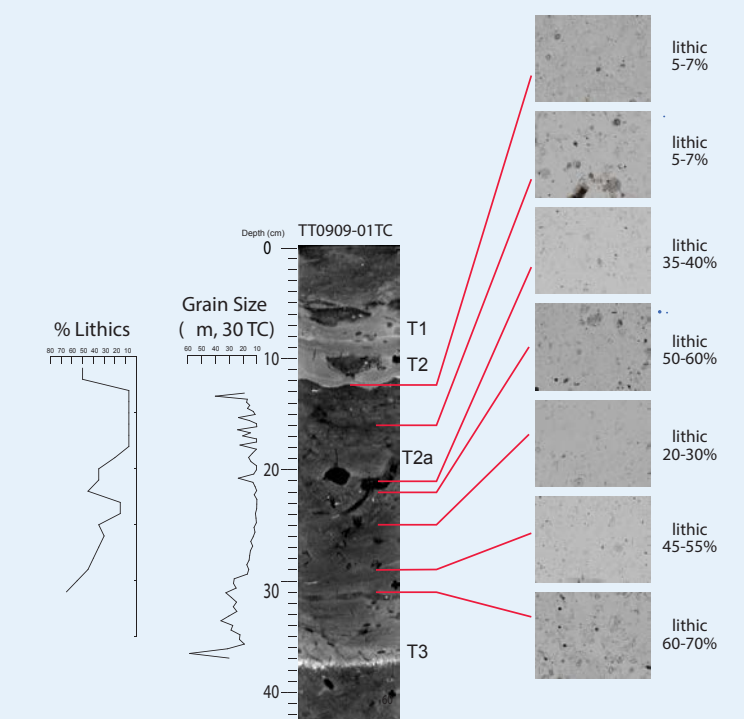
Multiple fining-upward Bouma T_{A-C} sequences, capped by a fine-grained fining-upward tail associated with the waning turbidity current (Bouma T_D).



Physical properties (density, magnetic susceptibility) often increase with grain size and can serve as proxies

“Mud turbidites”:

Thinner mud-silt beds with more subtle grain size and color variations; increased bioturbation; similar structures to sandier turbidites; high lithic content



From Goldfinger et al., 2012

MARINE TURBIDITE RECORD: CORRELATION

STRATIGRAPHIC "FINGERPRINTING"

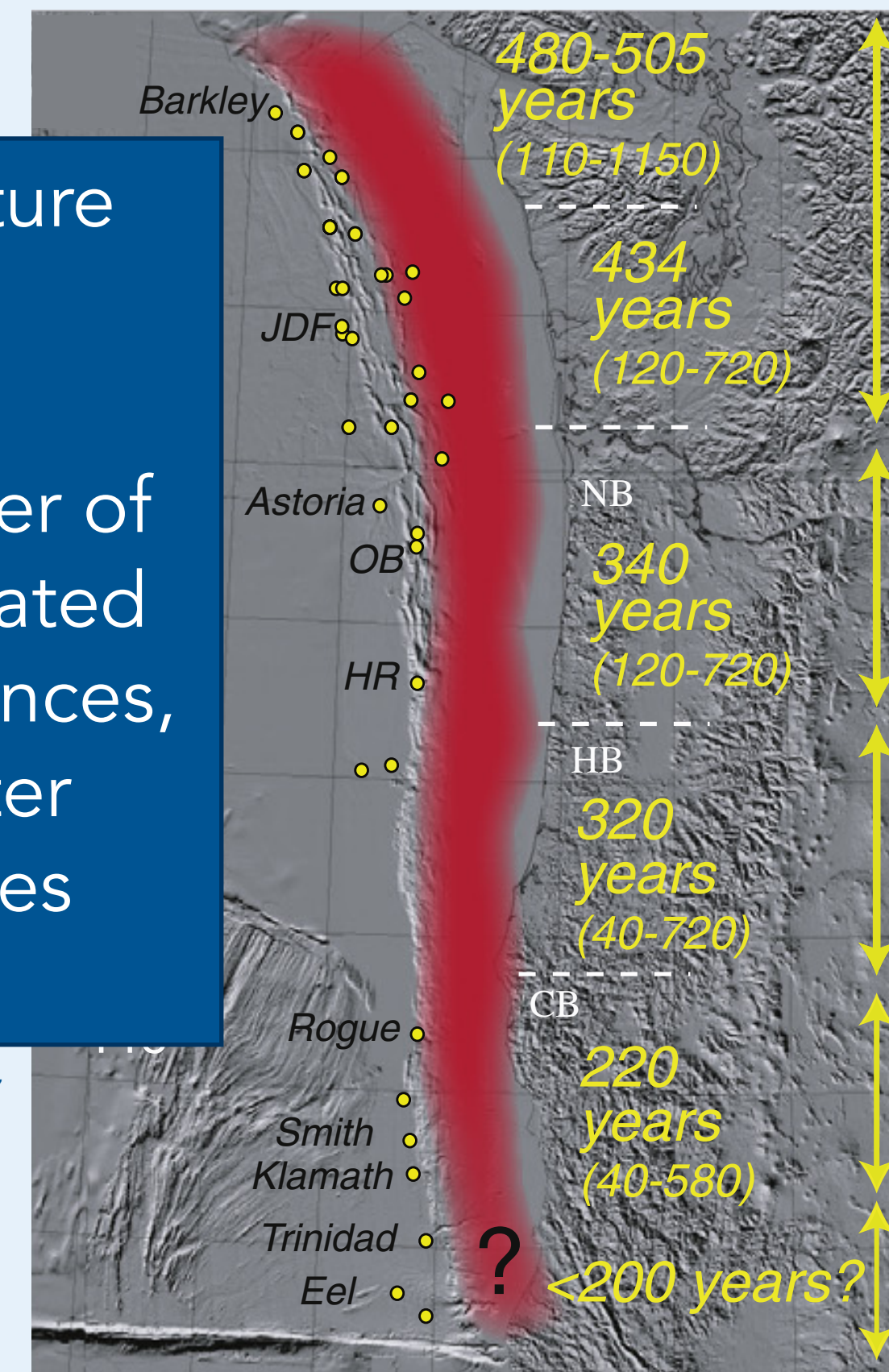
Stratigraphic correlation of turbidites within and between systems, over 100s of kms, relies primarily on the comparison of:

Physical properties (*Magnetic susceptibility and gamma density as proxies for grain size, supported by direct grain size subsample measurements, RGB color, P-wave velocity, CT and X-ray imagery*)

and is guided or bounded by age constraints from:

Marker beds (*Pleistocene/Holocene; Mazama Ash datum*) and **Radiometric dating methods** (^{14}C ; ^{210}Pb)

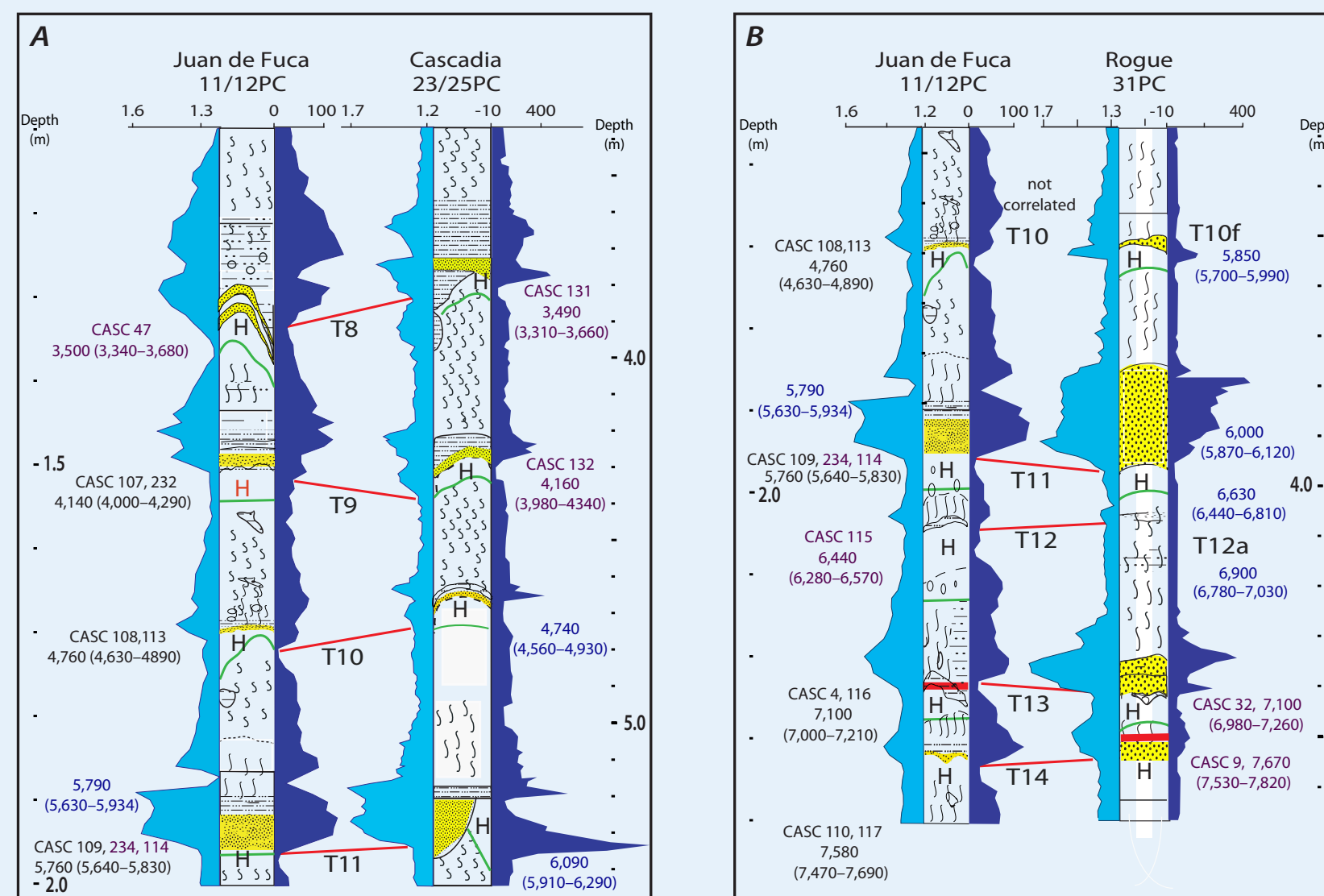
Full margin rupture recurrence ~500 yrs;
Increasing number of turbidites, correlated over shorter distances, suggests shorter recurrence times in the south



Goldfinger et al., 2017

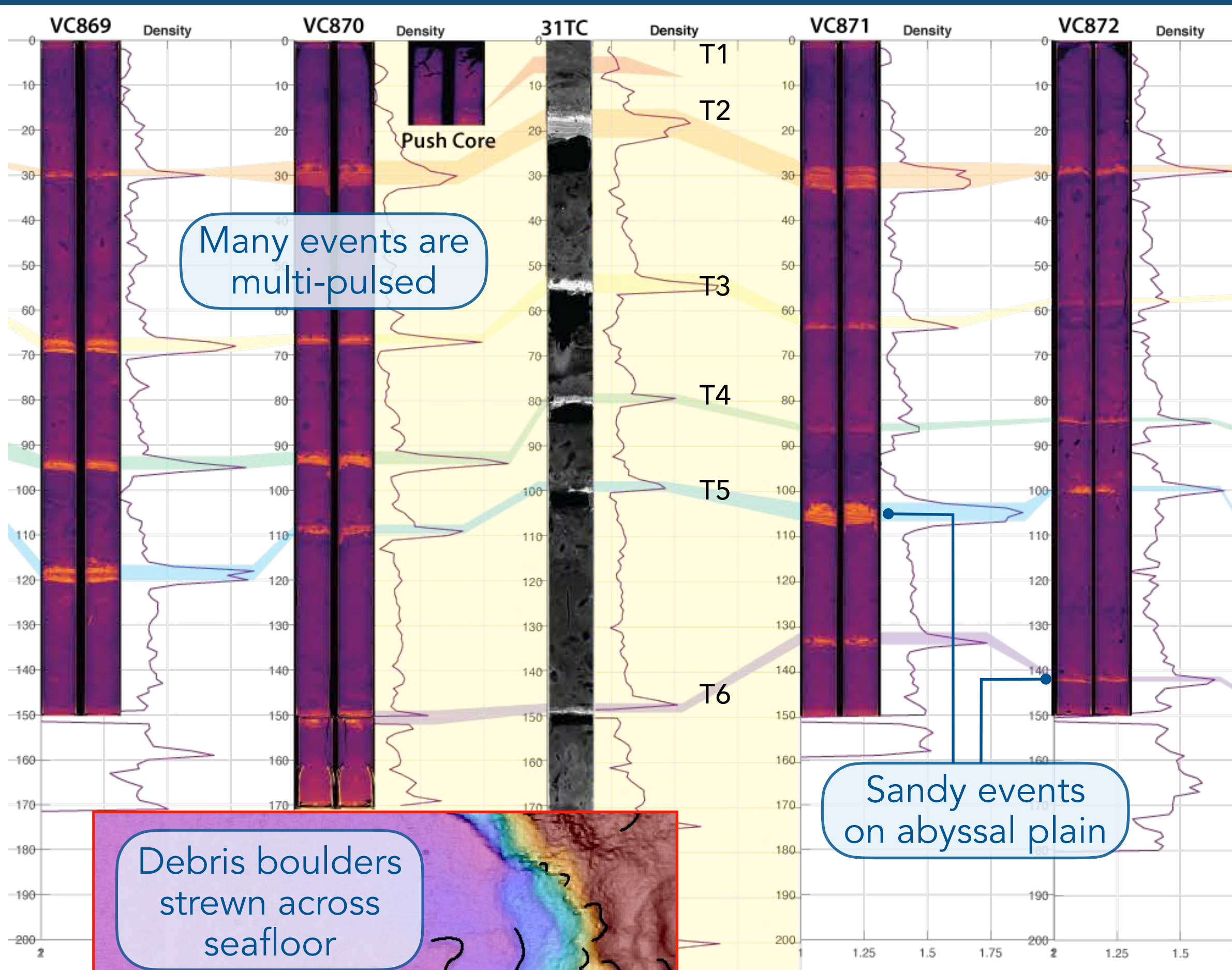
EXPLANATION	LITHOLOGY
CASC 11: 7,290 (7,220-7,380) Sample # AMS ^{14}C age and 2σ range	Hemipelagic clay
CASC 11: 7,290 (7,220-7,380) Erosion corrected AMS ^{14}C age and 2σ range	Turbidite silty mud
7,290 (7,220-7,380) Hemipelagic age and estimated 2σ range	Silt
CASC 11: 7,290 (7,220-7,380) Reversed AMS ^{14}C age and 2σ range	Very fine sand
Oldest Mazama ash bearing turbidite	Sand
Radiocarbon sample location	Mottled clay
Core break	Burrows
High-resolution point magnetics	Pleistocene
Density (g/cm^3)	Shell
Hemipelagic basal boundary	Wood fragment
Correlation lines, dashed if uncertain	
H/P Holocene/Pleistocene boundary	

From Goldfinger et al., 2012

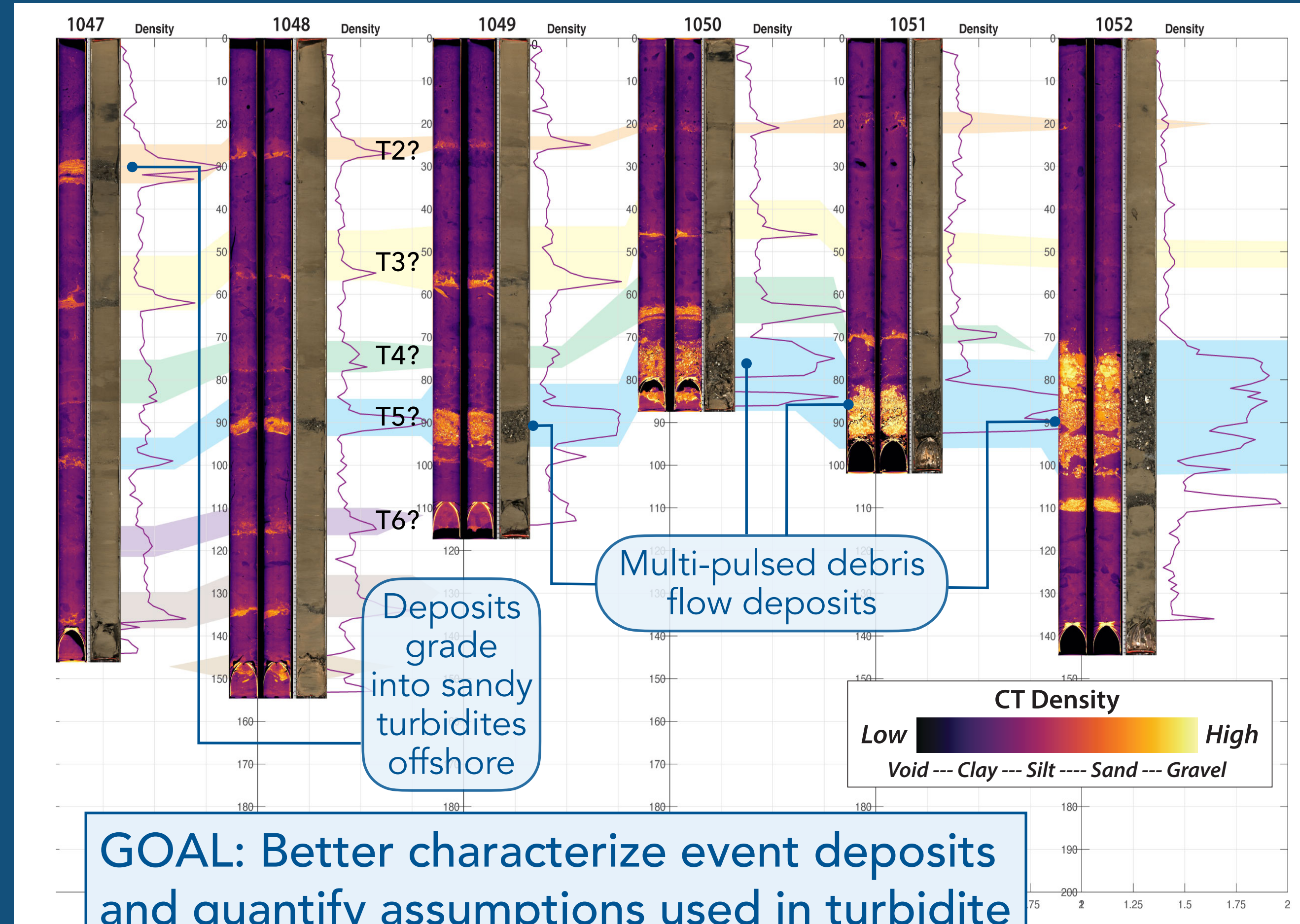


Uncertainties (ages, robustness of long distance correlation, interpretation of deposits) make it difficult to distinguish the length of ruptures, as well as the spatial extent and magnitude of shaking — requires better understanding of turbidite generating systems in Cascadia

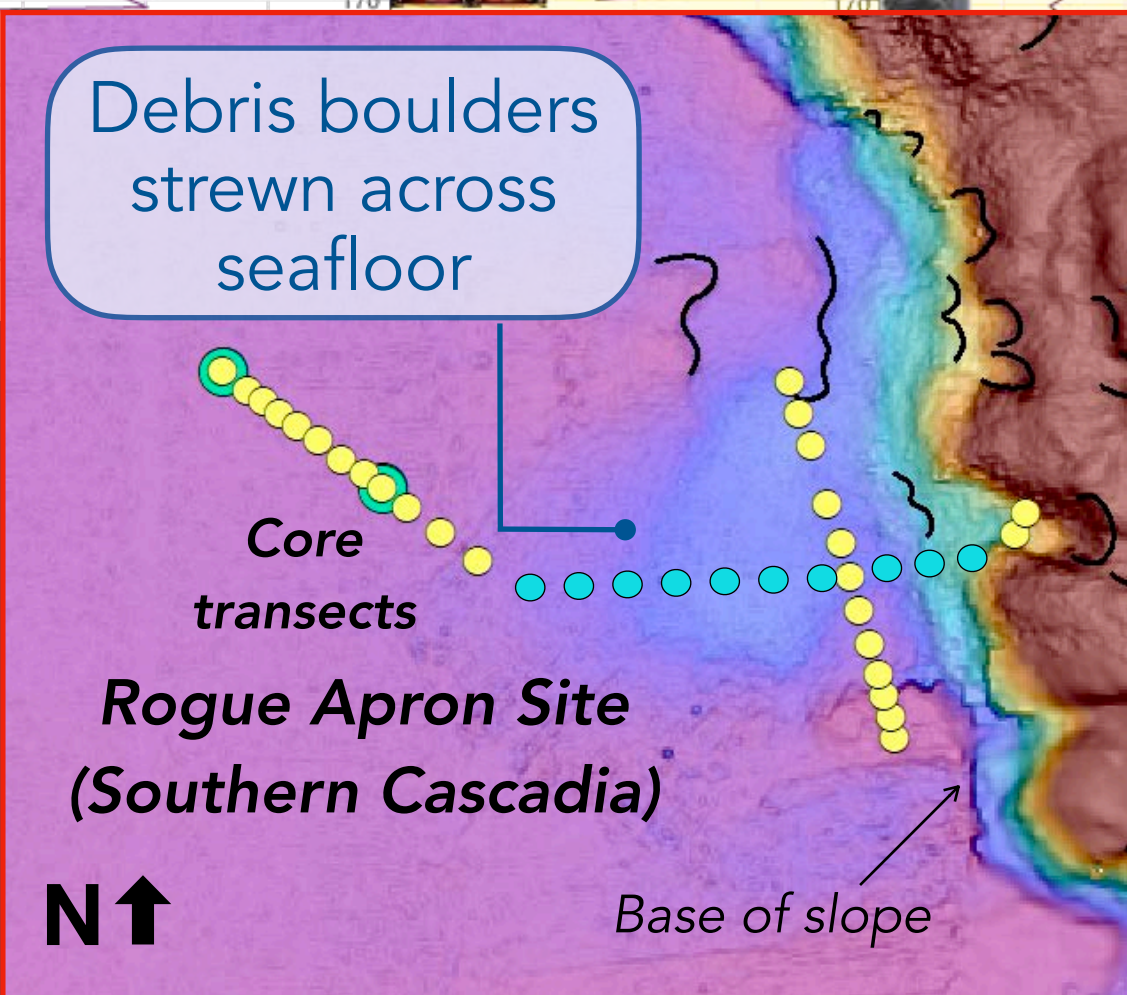
VARIABILITY WITHIN TURBIDITE SYSTEMS



31TC and T-events from Goldfinger et al., 2012



GOAL: Better characterize event deposits and quantify assumptions used in turbidite age models to reduce uncertainties

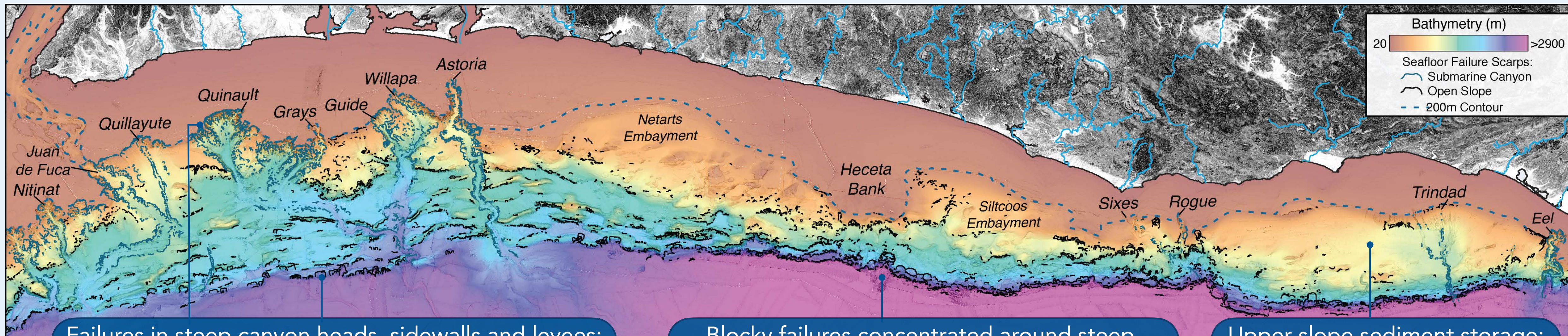


Vibracore transects (200-400m spacing), paired with push cores to collect undisturbed sediment-water interface, and ROV video observations, show *significant variability both between and within turbidite systems*

TURBIDITE SOURCES: CANYONS VS. OPEN SLOPE

What are the sources and pathways of turbidity flows across Cascadia?

← Northern → ← Central → ← Southern →



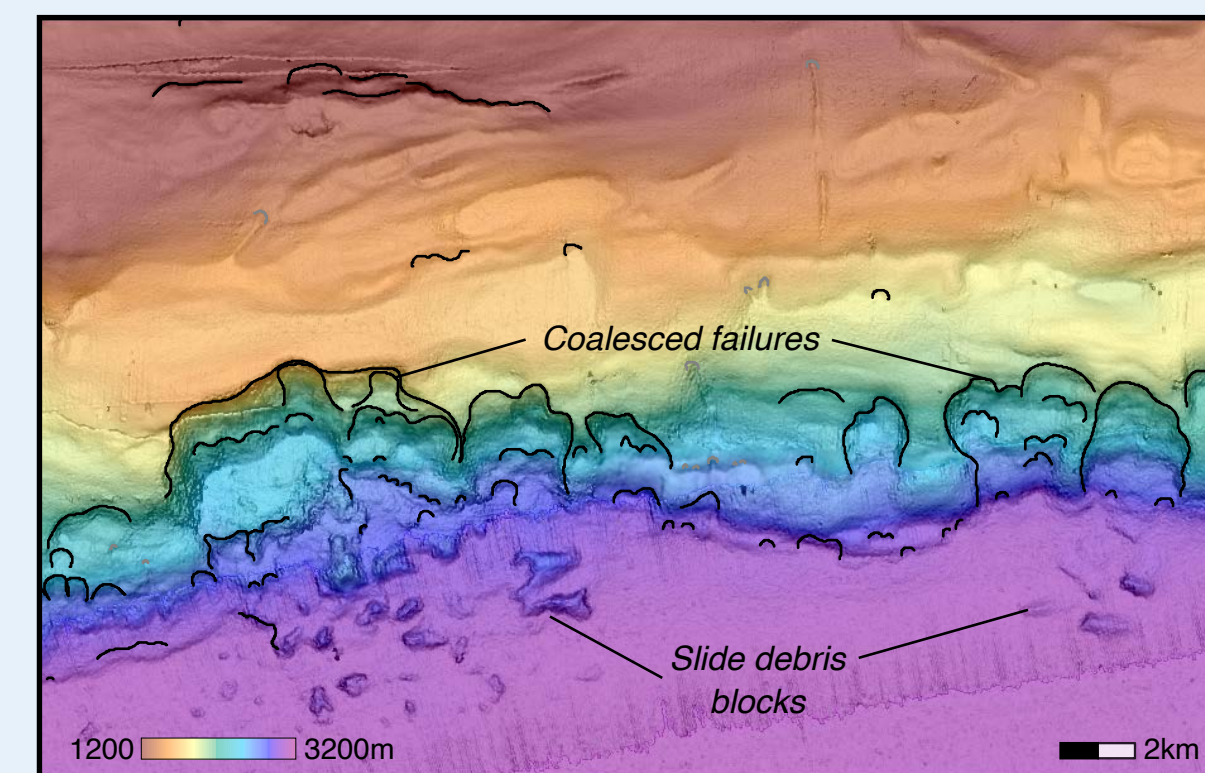
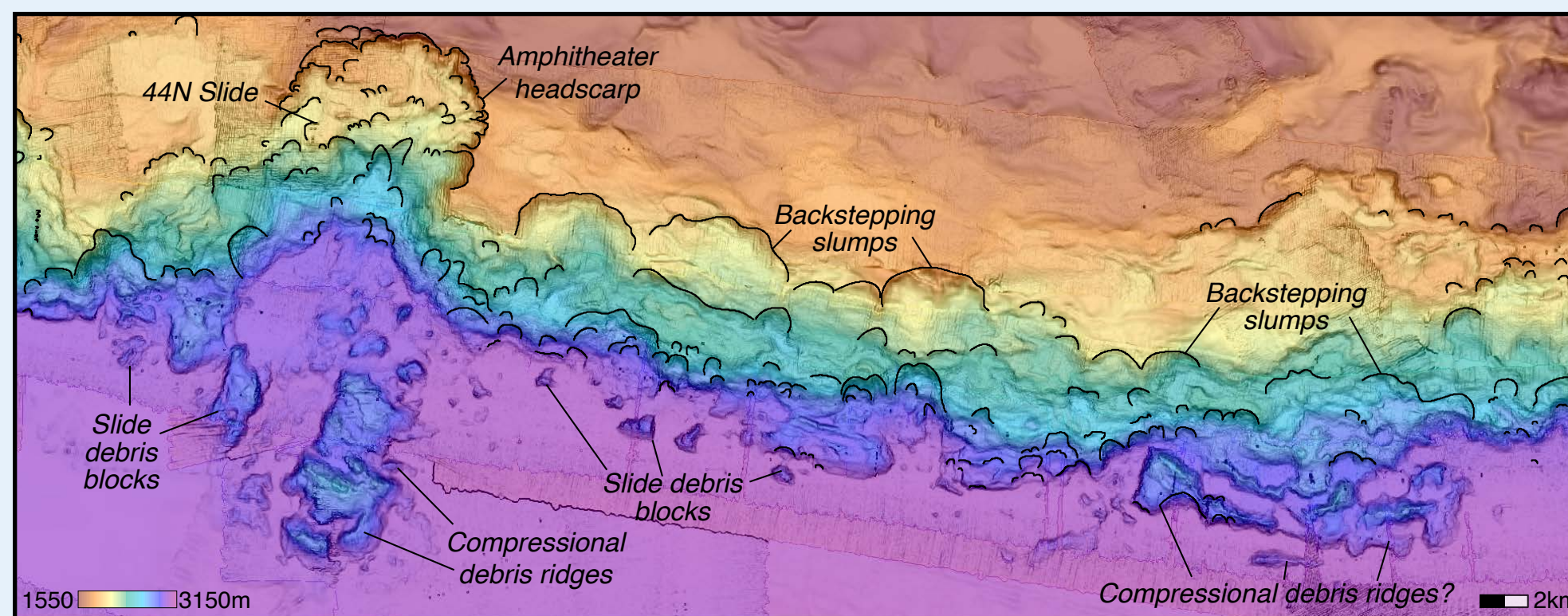
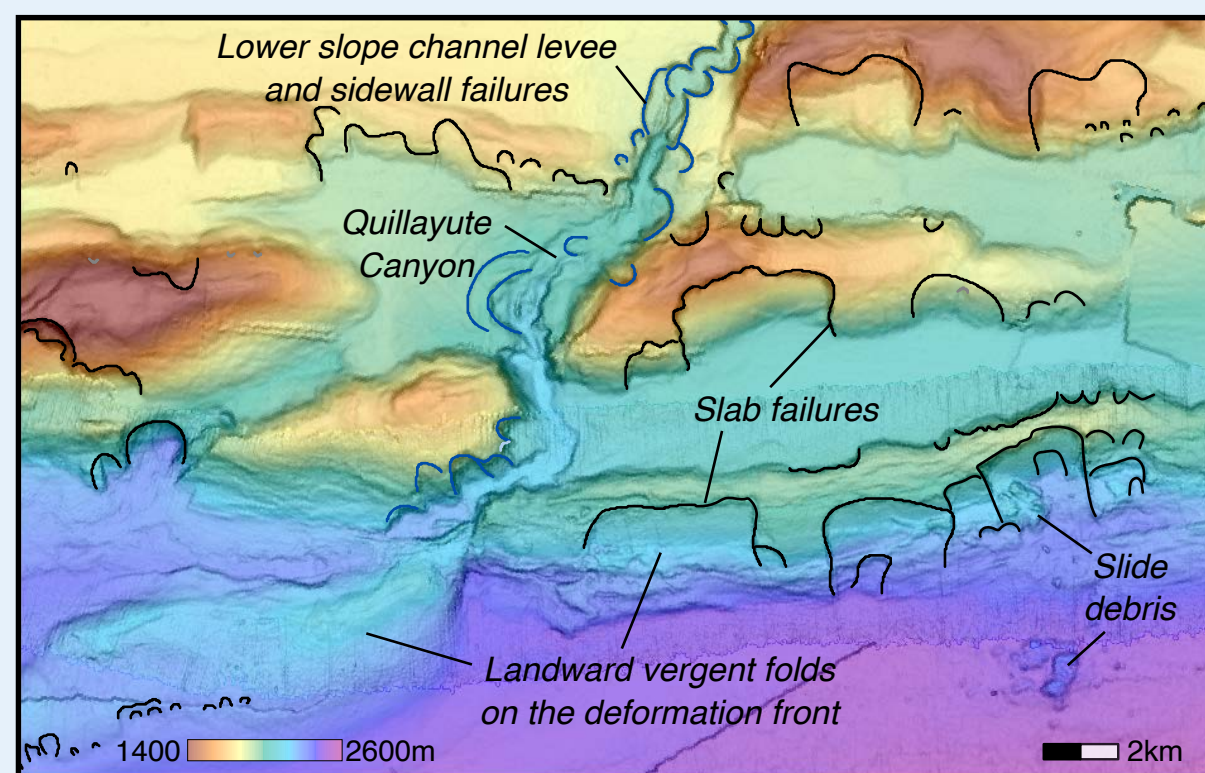
Failures in steep canyon heads, sidewalls and levees;
On the steep faces of landward vergent folds

Blocky failures concentrated around steep,
stepped terrace morphology of the lower slope

Upper slope sediment storage;
Failures on steep lower slope

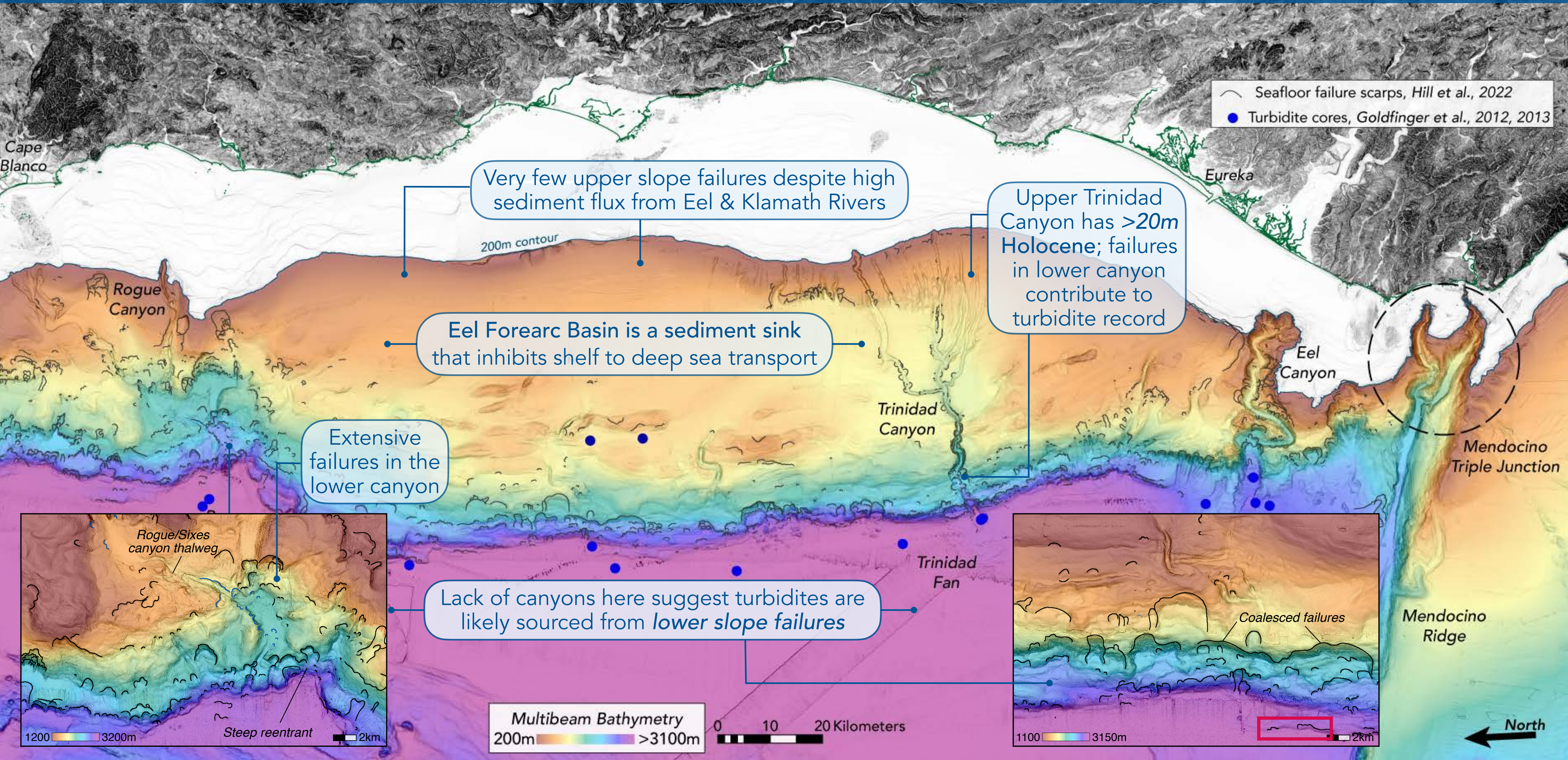
Sediment storage on the upper slope and pervasive mass wasting along the steep lower slope suggests *disintegration of the outer wedge is the primary source of turbidites in Cascadia*

Hill et al.,
EPSL, 2022

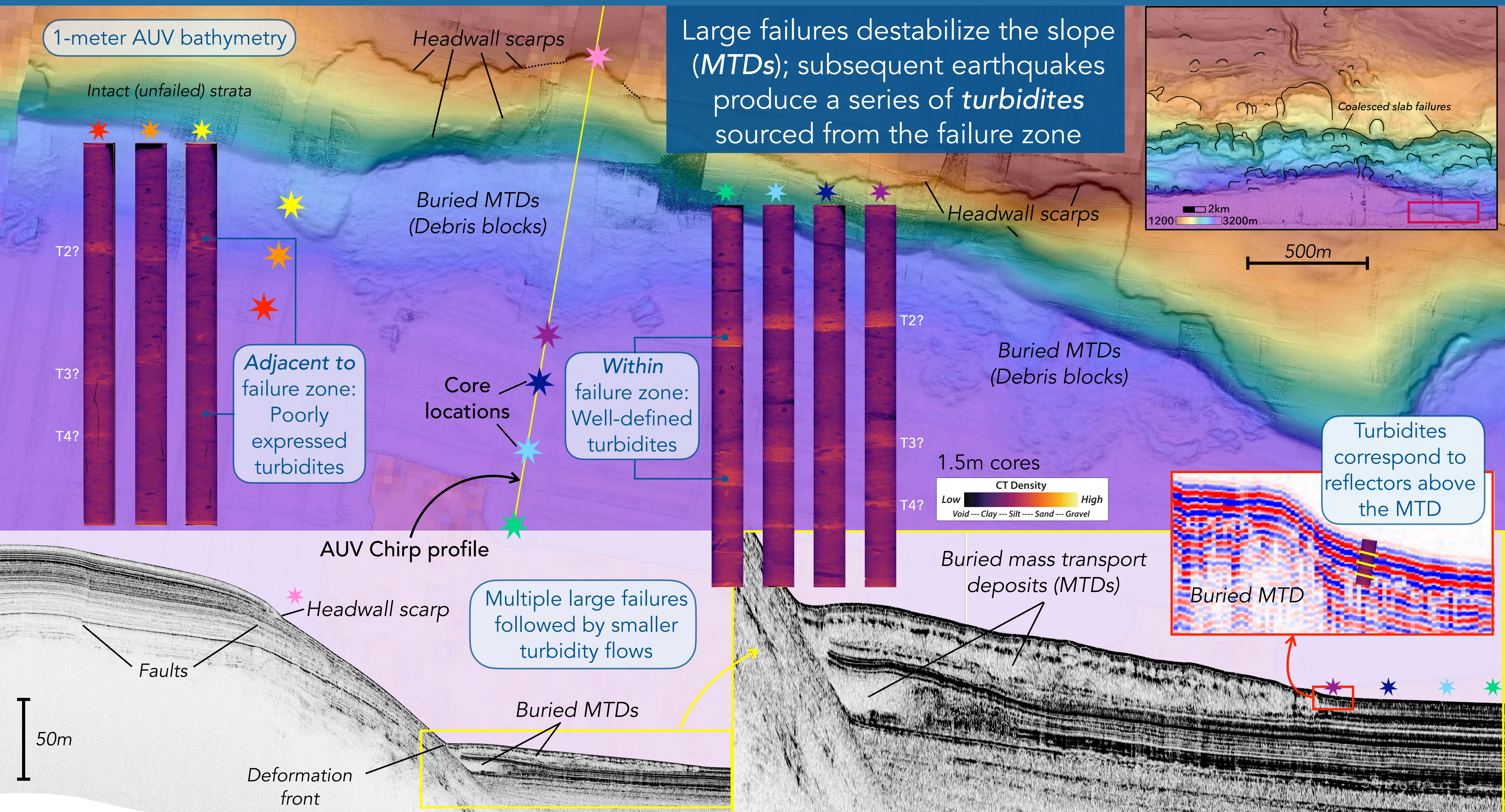


SOUTHERN CASCADIA: LOWER SLOPE FAILURES

Sediment storage on the upper slope and pervasive mass wasting along the steep lower slope

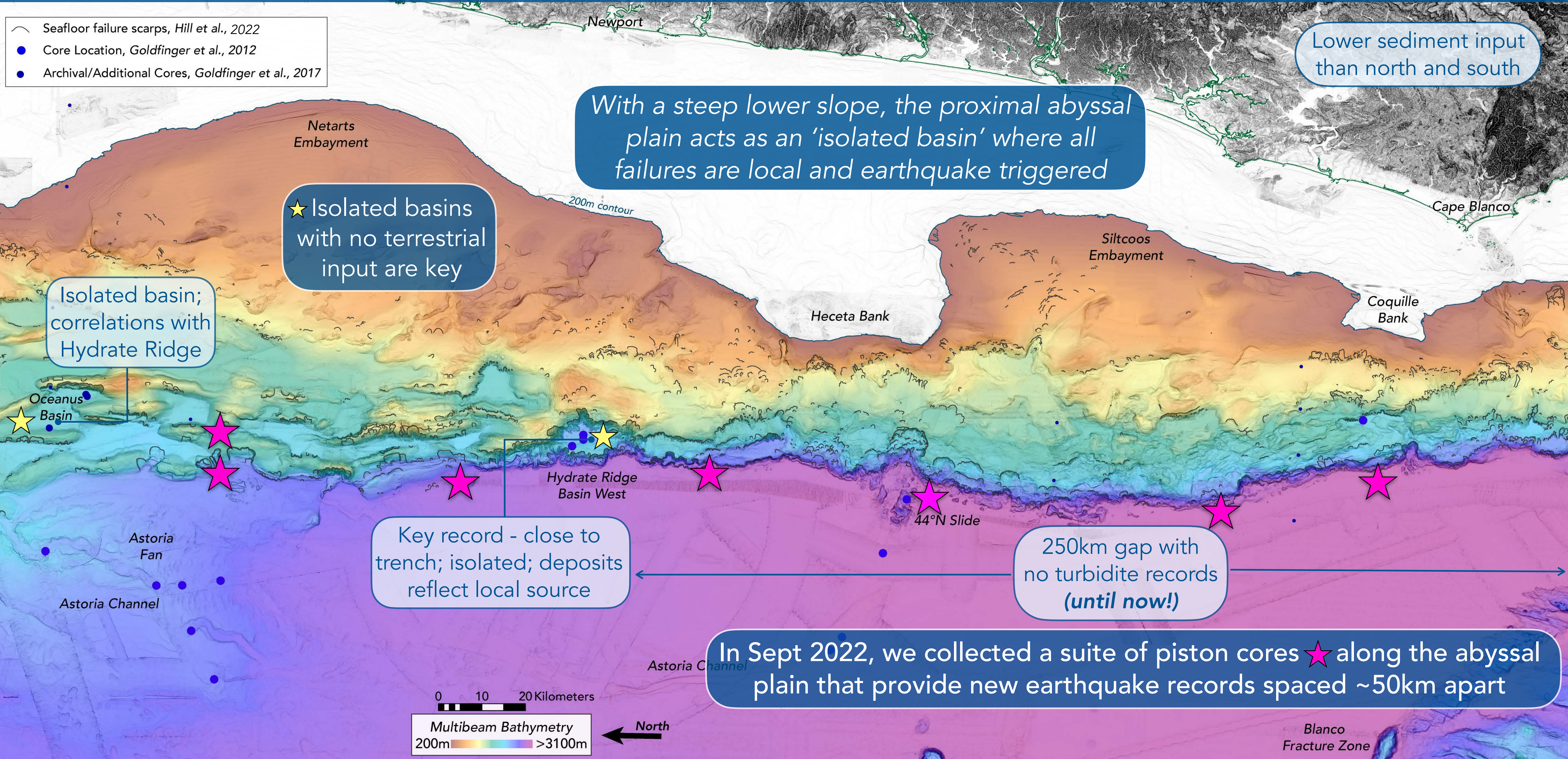


Earthquake generated turbidites appear to be sourced from seafloor failures on the lower slope in southern Cascadia



CENTRAL CASCADIA: NO SUBMARINE CANYONS

Steep lower slope with stepped terrace morphology and extensive failure zones



- Seafloor failure scarps, Hill et al., 2022
- Core Location, Goldfinger et al., 2012
- Archival/Additional Cores, Goldfinger et al., 2017

Lower sediment input than north and south

With a steep lower slope, the proximal abyssal plain acts as an 'isolated basin' where all failures are local and earthquake triggered

★ Isolated basins with no terrestrial input are key

Isolated basin; correlations with Hydrate Ridge

Oceanus Basin

Hydrate Ridge Basin West

44°N Slide

Key record - close to trench; isolated; deposits reflect local source

250km gap with no turbidite records (until now!)

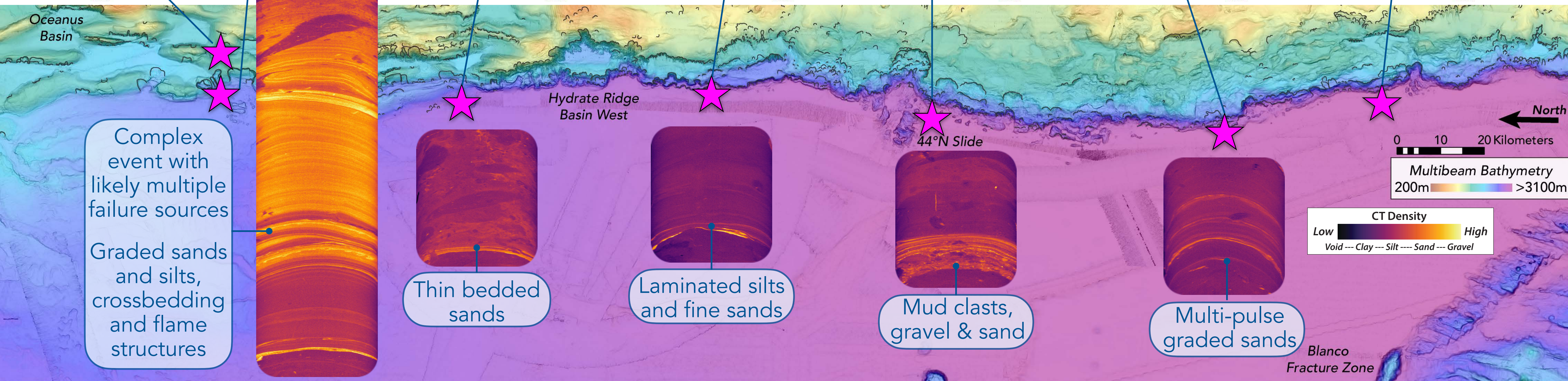
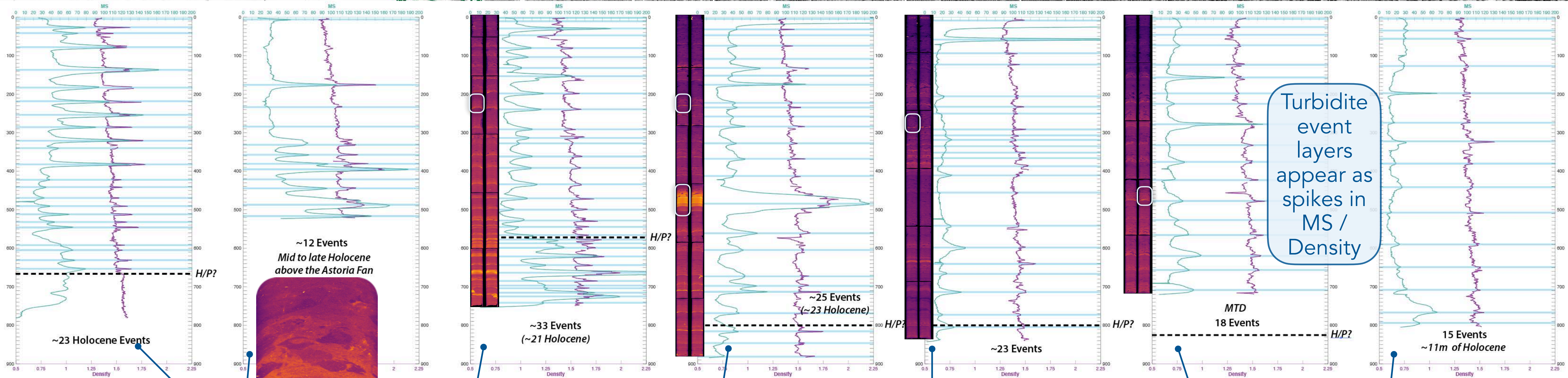
In Sept 2022, we collected a suite of piston cores ★ along the abyssal plain that provide new earthquake records spaced ~50km apart



Blanco Fracture Zone

CENTRAL CASCADIA: ABYSSAL TURBIDITE RECORDS

Initial estimates suggest 20-23 Holocene events likely sourced from proximal failures of the lower slope



Complex event with likely multiple failure sources
Graded sands and silts, crossbedding and flame structures

Thin bedded sands

Laminated silts and fine sands

Mud clasts, gravel & sand

Multi-pulse graded sands

Turbidite event layers appear as spikes in MS / Density

Multibeam Bathymetry
200m -> >3100m

CT Density
Low -> High
Void --- Clay --- Silt --- Sand --- Gravel

Oceanus Basin

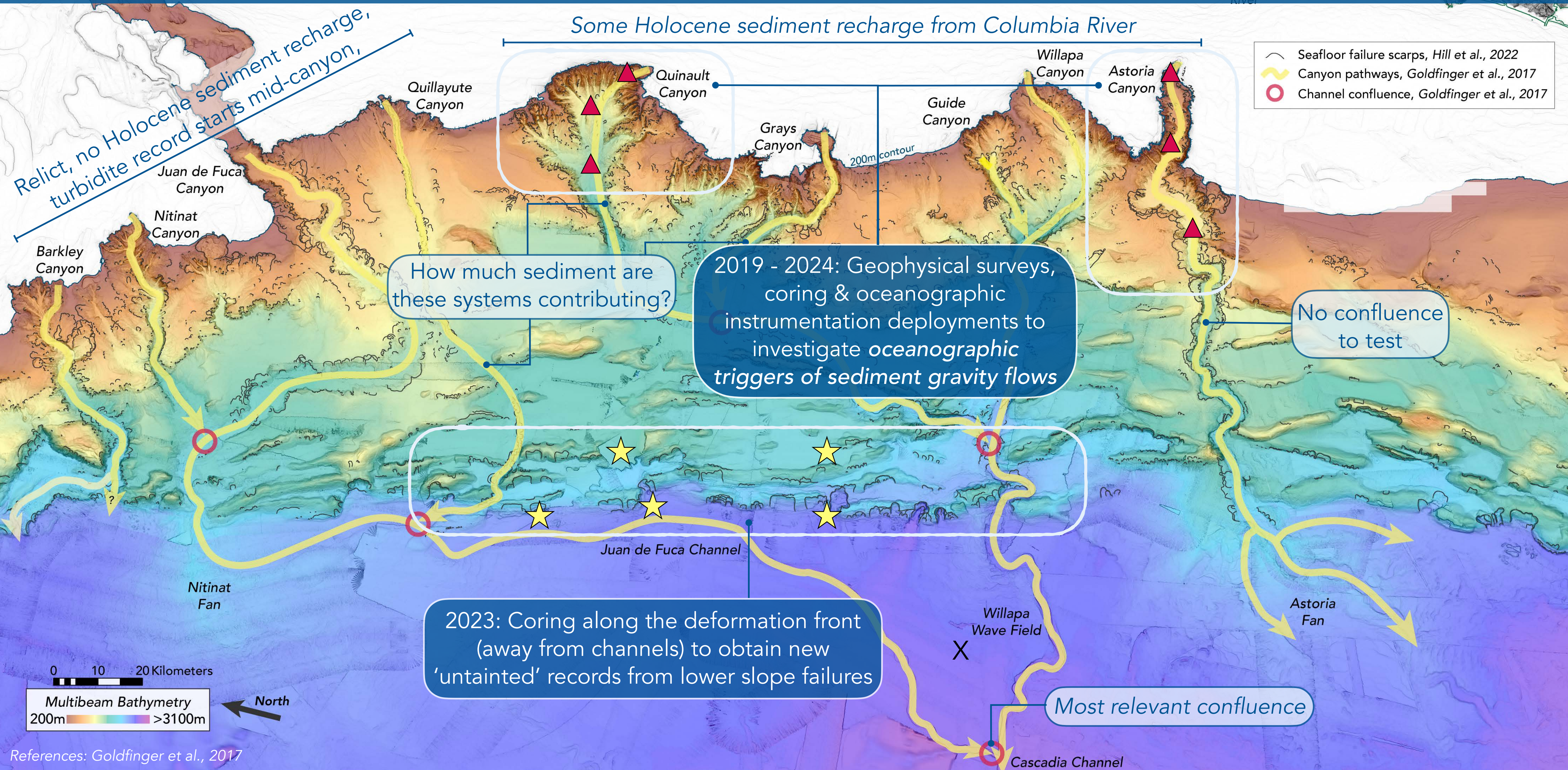
Hydrate Ridge Basin West

44°N Slide

Blanco Fracture Zone

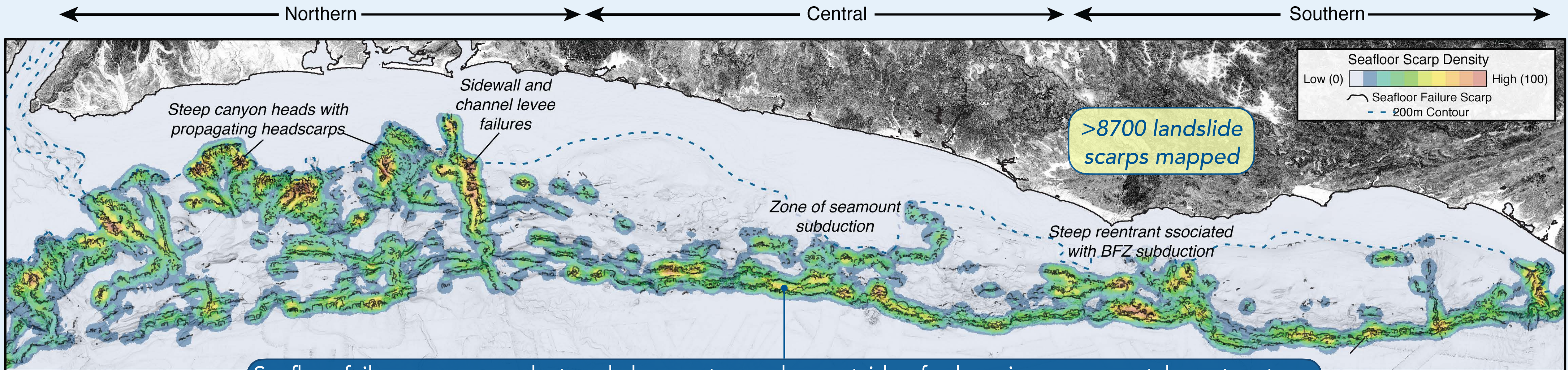
NORTHERN CASCADIA: WHAT IS THE ROLE OF CANYONS?

Well developed canyons with limited Holocene sediment input & substantial lower slope inputs form sidewall failures



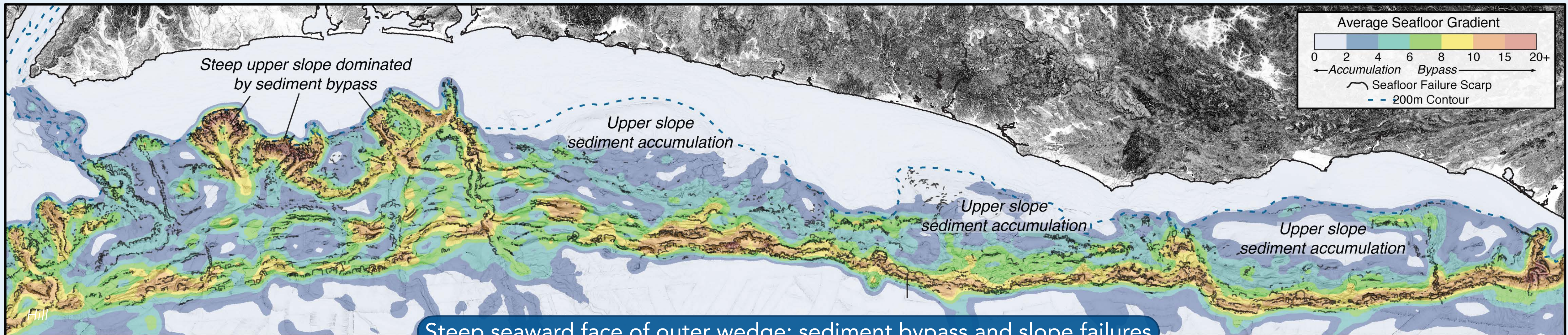
TURBIDITE SOURCES: LOWER SLOPE FAILURES

Earthquake generated turbidites appear to be sourced from seafloor failures on the lower slope in all regions of Cascadia



Seafloor failure scarps are clustered along outer wedge, outside of submarine canyon catchment systems

From: Hill et al., EPSL, 2022



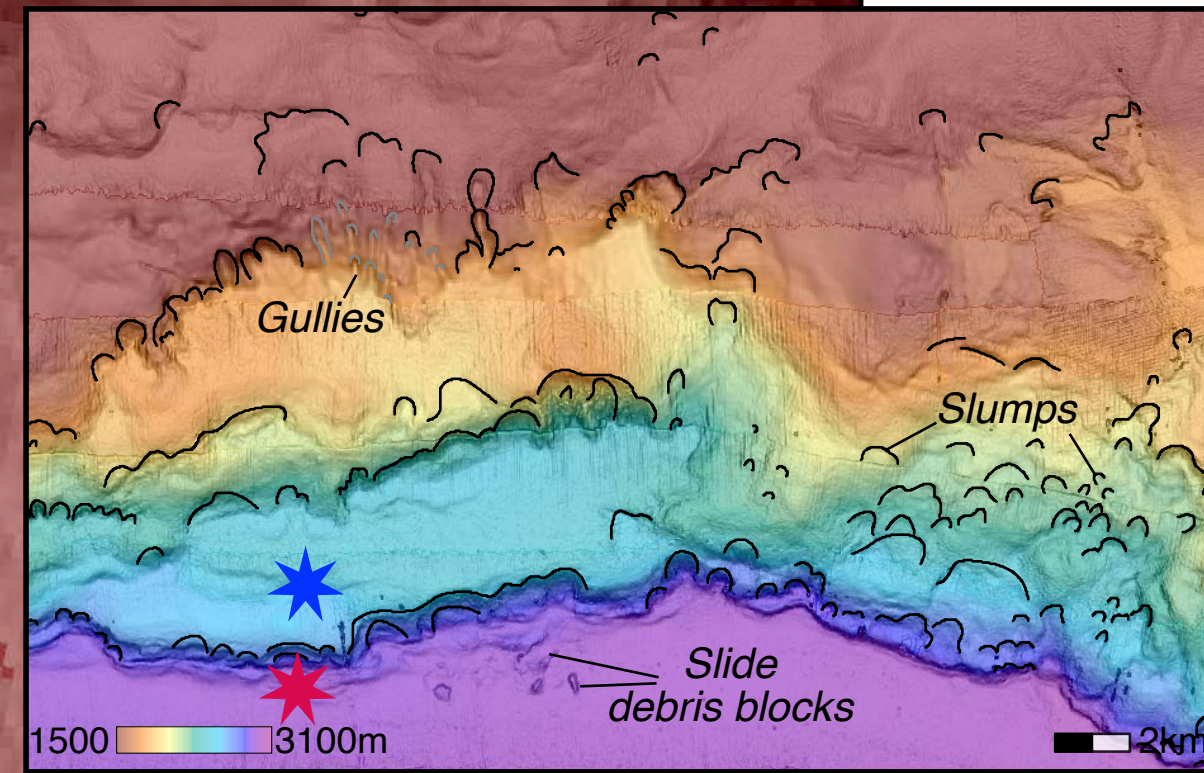
Steep seaward face of outer wedge: sediment bypass and slope failures

TECTONIC OVERSTEEPENING & SLOPE FAILURE

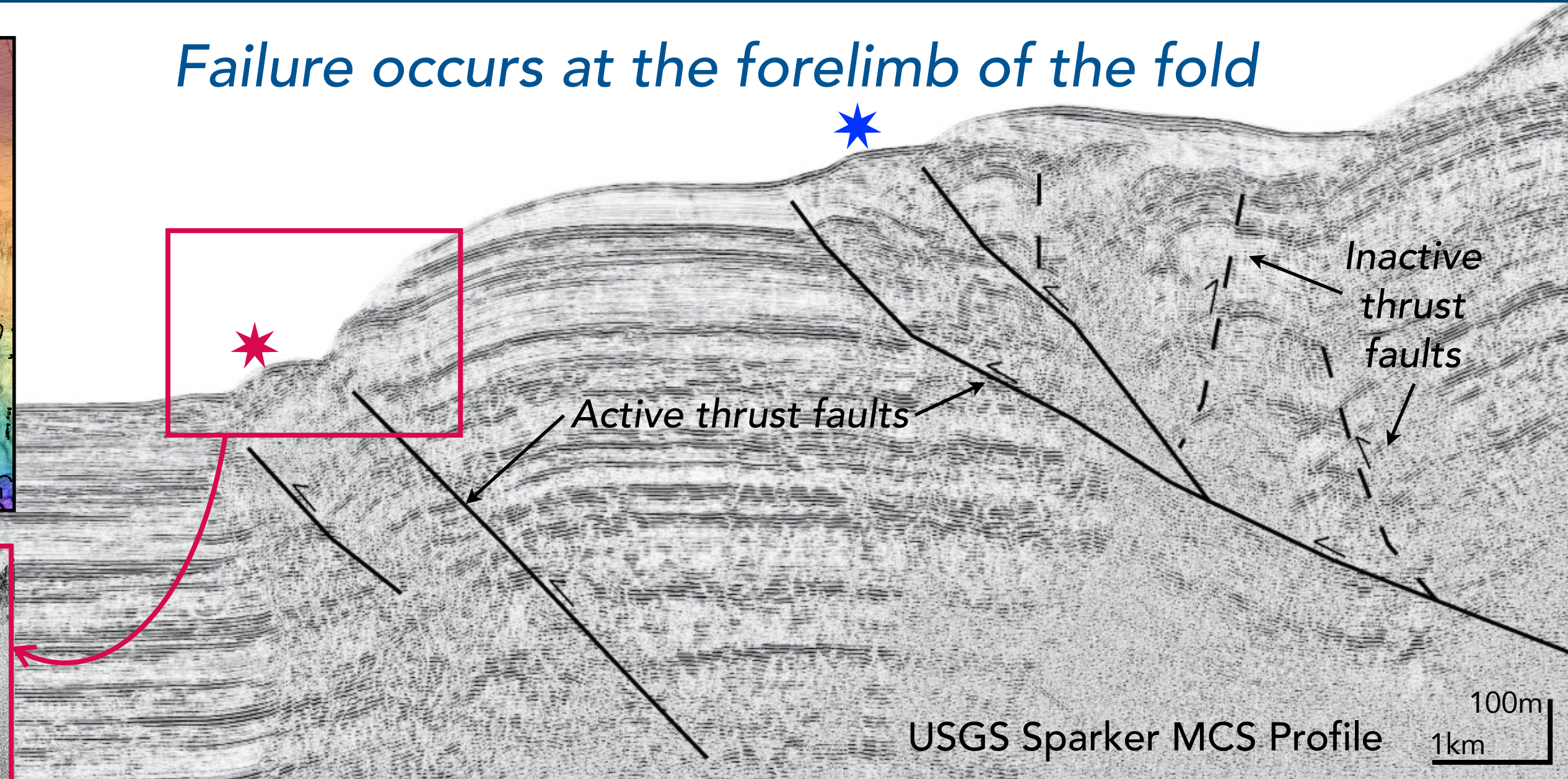
Imbricate thrust faults create a stepped terrace morphology with steep ledges in Central Cascadia

Failures are pervasive along the deformation front — at all scales

1-meter AUV bathymetry



Failure occurs at the forelimb of the fold



Surficial failures

Steep scarps with outcropping bedding

Shallow tension cracks

Steep failure planes with truncated strata

Deformation front

AUV Chirp Profile

Uplift and deformation of the outermost wedge leads to recurring, earthquake-triggered failures that are recorded in the abyssal turbidite record

Steep headscarps

Landslide debris

Landslide debris blocks

500m

MCS profile

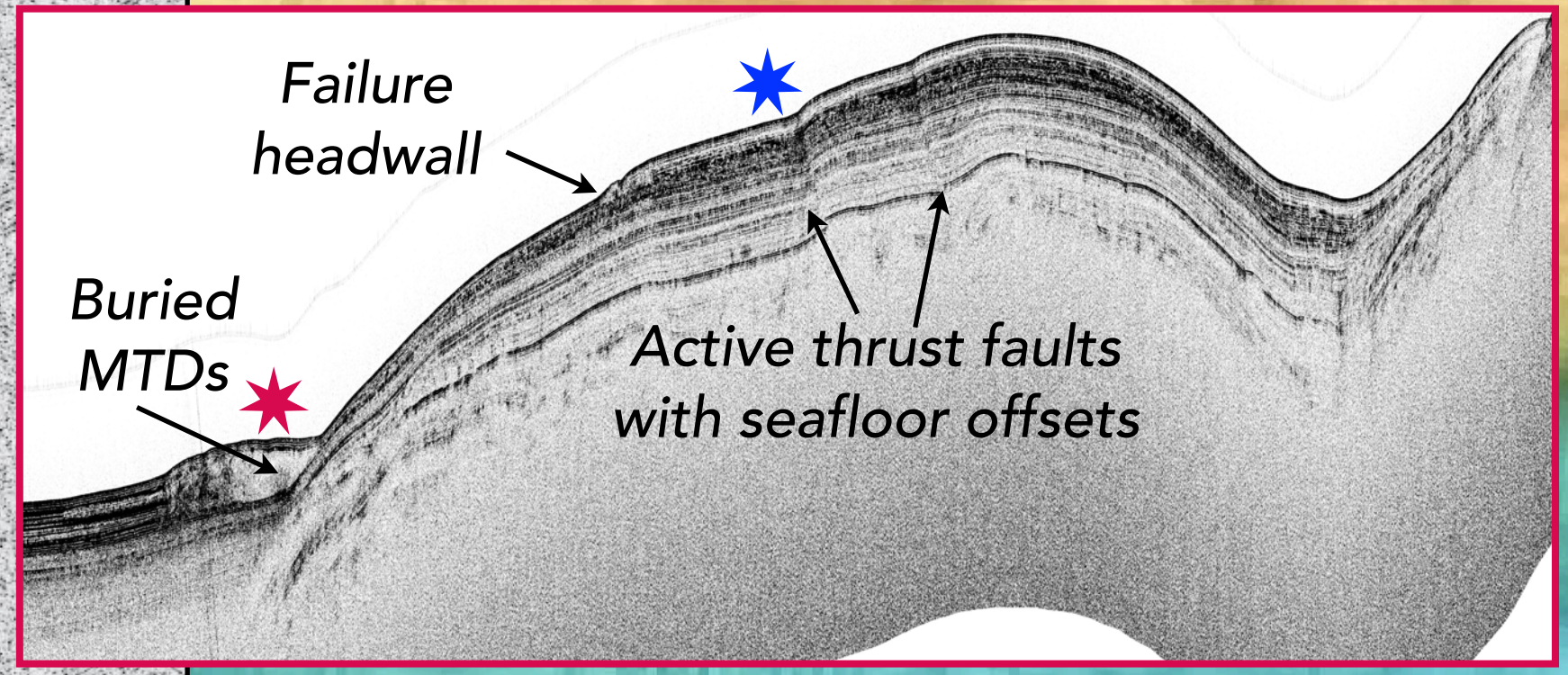
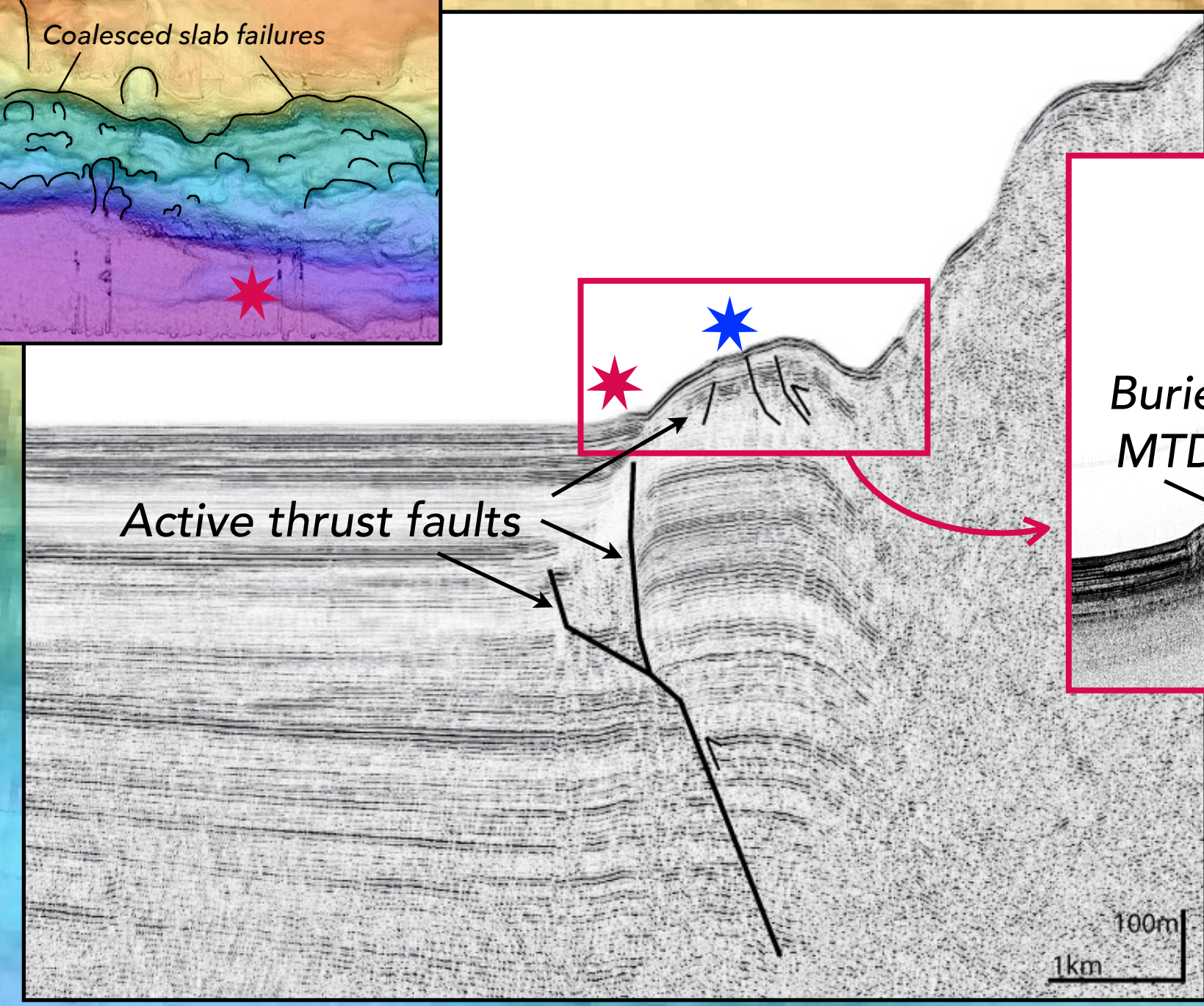
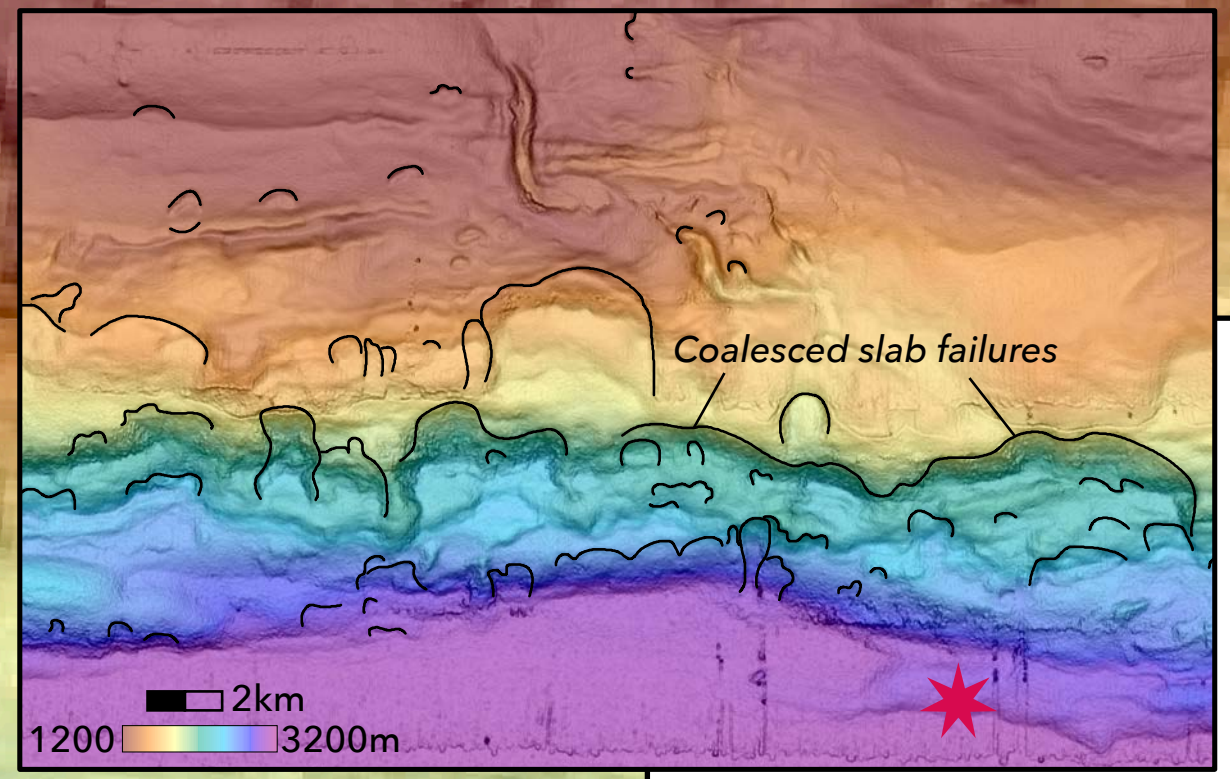
TECTONIC OVERSTEEPENING & SLOPE FAILURE

Steeply dipping thrust faults create lower relief folds in southern Cascadia

MCS profile

1-meter AUV bathymetry

Uplift with every earthquake cycle creates oversteepening that outpaces the effects of compaction strengthening and preconditions the slope for failure



Failures occur at wide variety of slope gradients; average slope here: 4-6°

Accretion of abyssal sediment into the outer wedge provides unlimited recharge of sediment to produce abyssal seismoturbidites during slope failure

Seafloor offset thrust faults

Failure headwall

Landslide debris

500m

SUMMARY

The abyssal turbidite record is a GREAT source of earthquake triggered event deposits.

Local sources are best!



Abyssal turbidites sourced from lower slope failures (away from canyons and channels) avoid many of the pitfalls and arguments commonly made against turbidite stratigraphy:

- (1) Many abyssal turbidites can be tied to mass transport deposits at the base of the slope — the only viable source for these abyssal turbidites is earthquake triggered ground failure
- (2) Correlation should not be solely based on matching physical properties or counting the number of events — it is OK if events are missing or look different in different places — they are probably sourced locally
- (3) Tectonic oversteepening along the lowermost slope provides infinite recharge of sediment to fail during earthquake shaking

Better understanding of turbidite generating systems and refined chronologies from detailed core transects are key to interpreting the offshore record.

*We need to better constrain the minimum threshold for shaking induced failure
The spatial distribution of landslides can inform where/how shaking occurs*