

Quaternary Fault and Fold Database of the United States

As of January 12, 2017, the USGS maintains a limited number of metadata fields that characterize the Quaternary faults and folds of the United States. For the most up-to-date information, please refer to the [interactive fault map](#).

Wallula fault system (Class A) No. 846

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Synopsis	The Wallula fault system is a prominent northwest-striking fault zone that extend from near Milton-Freewater, Ore., to near Kennewick, Wash., including the fault adjacent to The Rattles. This fault zone is collocated with several regional-scale lineaments, including the Cle Elum-Wallula deformed zone (CLEW), the Rattles Wallula trend, alignment, or lineament (RAW), and the Olympic-Wallowa lineam (OWL). In the northwest, movement on the fault system results in anticlinal folds up to 250 m of structural relief (Coppersmith and others, 2014 #7402) in rocks of Miocene Columbia River Basalt Group. The Wallula fault system is variously ma as linear, steeply dipping strike-slip, normal, or reverse faults in Quaternary surfic deposits and rocks of the Miocene Columbia River Basalt Group. The mapped fa pattern, local observations of subhorizontal slickensides, and north-south compre stress regime support a right-lateral strike slip sense of slip on the Wallula fault. T age of most-recent faulting in the Wallula fault system is poorly known; several s suggest latest Quaternary displacement at least one location.
Name	The Wallula fault system is a prominent northwest-striking fault zone that include

comments	<p>several named faults in Oregon: the Bade, Barrett, Dry Creek, Forks, Little Dry C Milton-Freewater, Pine Creek, Promontory Point, Umapine, Wallula, and Wallula faults (Kienle and others, 1979 #3728; Mann and Meyer, 1993 #3535; McQuarrie 1993 #4337; Reidel and others, 1994 #3539).</p> <p>Fault ID: These structures are included in fault number 11 of Pezzopane (1993 # and fault number 77 of Geomatrix Consultants, Inc. (1995 #3593).</p>
County(s) and State(s)	<p>BENTON COUNTY, WASHINGTON WALLA WALLA COUNTY, WASHINGTON UMATILLA COUNTY, OREGON</p>
Physiographic province(s)	<p>COLUMBIA PLATEAU</p>
Reliability of location	<p>Good Compiled at 1:100,000 scale.</p> <p><i>Comments:</i> Location of fault in Oregon from ORActiveFaults (http://www.oregongeology.org/arcgis/rest/services/Public/ORActiveFaults/Map5 downloaded 06/02/2016) attributed to 1:100,000-scale mapping of Maden and Ge (2007 #7804). Location of fault in Washington from GER_Seismogenic_WGS84 (http://www.dnr.wa.gov/publications/ger_portal_seismogenic_features.zip, downloaded 05/23/2016) attributed to 1:100,000-scale maps of Schuster 1994 #4 and Reidel and Fecht (1994 #5565). Magnetic surveys show that the Wallula fault located 350 m to the southwest of the trace shown on published maps, passes directly through deformed late Pleistocene or younger deposits exposed at Finley quarry, extends uninterrupted over 120 km; Sherrod and others (2016 #7690) illustrate the mislocation of the fault in figure 4 at detail greater than the metadata mapping accuracy.</p>
Geologic setting	<p>The Wallula fault system is a prominent northwest-striking fault zone in the southeastern part of the Yakima fold belt and appears to link with northwest-striking folds and thrust faults of the Rattlesnake Hills structures [565] in the Rattlesnake directly to the north. The Wallula fault system is included in several regional-scale lineaments, including the Cle Elum-Wallula deformed zone (CLEW), the Rattlesnake Wallula trend, alignment, or lineament (RAW), and the Olympic-Wallowa lineament (OWL) (Reidel and others, 1994 #3539). The Yakima fold belt consists of anticlinal folds and thrust faults in Miocene Columbia River Basalt Group. Strain across the Yakima fold belt at this latitude appears to be locally confined to the narrow Wallula fault zone south of the Pasco gravity low (Blakely and others, 2014 #7406). The Wallula fault system is mostly mapped as linear, steeply dipping strike-slip, normal reverse faults in Quaternary surficial deposits and rocks of the Miocene Columbia River Basalt Group (Kienle and others, 1979 #3728; Swanson and others, 1981 # Tolan and Reidel, 1989 #3765; Walker and MacLeod, 1991 #3646; Hutter, 1997 #5650). Coppersmith and others (2014 #7402) estimate average structural relief a</p>

	the Wallula structures to be 250 m with a maximum of 300 m.
Length (km)	63 km.
Average strike	N53°W
Sense of movement	<p>Right lateral, Reverse</p> <p><i>Comments:</i> The sense and amount of displacement across this fault is poorly known. The northwestern continuation of the Wallula fault system in Washington consists of anticlinal folds and thrust faults, but the fault system is mostly mapped as linear, steeply dipping strike-slip, normal, or reverse faults in Oregon (Kienle and others, 1979 #3728; Swanson and others, 1981 #3496; Tolan and Reidel, 1989 #3765; Wagoner and MacLeod, 1991 #3646; Hutter, 1997 #5650). The mapped fault pattern, local observations of subhorizontal slickensides, and north-south compressive stress results support a right-lateral strike slip sense of slip on the Wallula fault system (Kienle and others, 1979 #3728; U.S. Army Corps of Engineers, 1983 #3480; Mann and Meyer, 1993 #3535; McQuarrie, 1993 #4337; Kuehn, 1995 #3478; 1996 #3530). However, Reidel and Tolan (1994 #3536) and Hutter and others (1994 #3525) interpret observations to infer that post-Miocene lateral displacement, if any, must be low. Pratt (2012 #3525) states that at least some right-lateral movement on the Wallula fault is required because of shortening in the Columbia Hills [568] and Horse Heaven Hills [567], in contrast to the lack of demonstrable shortening northeast of the Wallula fault. However, he notes that the modern stress field is less favorably oriented to produce strike-slip motion. Blum and others (2014 #7406) interpret magnetic anomalies identified in high-resolution aeromagnetic data that show clear evidence that Ice Harbor dikes are disrupted by the Wallula fault zone; they interpret the disruption to be a manifestation of right-lateral offset of individual dikes on the order of more than 1 km and total offset of 6.9 km in the past 8.5 m.y. Further interpretation of aeromagnetic and ground magnetic surveys suggest right lateral oblique (reverse) movement on the Wallula fault zone post emplacement of the 8.5-Ma Ice Harbor Member dikes, exhibiting up to 2.2 km of lateral displacement (Sherrod and others, 2016 #7690). In addition, Sherrod and others (2016 #7690) reports observing faint striae on the master fault that are subhorizontal suggesting reverse-dextral movement.</p>
Dip	<p>70–90°, V</p> <p><i>Comments:</i> Vertical to steeply northeast- and southwest-dipping fault planes have been observed along several faults in the Wallula fault system; a few gently-dipping fault planes have also been observed, but these are attributed to flower or "palm tree" fault patterns (Kienle and others, 1979 #3728; Mann and Meyer, 1993 #3535; McQuarrie, 1993 #4337; Kuehn, 1995 #3478; Hutter, 1997 #5650).</p>
Paleoseismology	Several early sites including Milton-Freewater quarry and Little Dry Creek (Kienle and others, 1979 #3728; Mann and Meyer, 1993 #3535; McQuarrie, 1993 #4337; Kuehn, 1995 #3478; Hutter, 1997 #5650).

studies

others, 1979 #3728), natural and man-made exposures (Rigby and Othberg, 1979 #3738; Farooqui and Thoms, 1980 #5824; Foundation Sciences Inc. 1980 #5722) three trenches near Yellepit Station Canyon (Gardner and others, 1981 #7691) do provide conclusive results regarding timing and the number of earthquakes. However, studies that addressed an exposure in Finley quarry near Wallula Gap on the Columbia River provide a variety of interpretations (Farooqui and Thoms, 1980 #5824; Foundation Sciences, Inc. 1980 #5722; Woodward-Clyde Consultants, 1981 #7690; Sherrod and others, 2016 #7690). Sherrod and others (2016 #7690) conclude the scarp formed from tectonic processes rather than flood erosion, and the colluvial deposits associated with the scarp. Coppersmith and others (2014 #7402) summarize the findings reported by Sherrod and others (2016 #7690) and other studies and conclude that Ice Age erosion and deposition might explain some features inferred to be the result of surface faulting at the Finley quarry; the Quaternary Studies Team favors earlier interpretations that there has been no surface deforming earthquakes that post date the deposition of deposits that are likely 71–58 ka. They do not completely dismiss the idea of younger faulting; they assign low weight to the possibility of a post 17±1.8 ka event based on the minimum age of unfaulted deposits based on $^{230}\text{Th}/\text{U}$ series dating and very low weight to the possibility of two or three faulting events that post date 17±1.8 ka.

Studies over the past few decades focused on an exposure in a quarry at Finley, WA (Farooqui and Thoms, 1980 #5824 Site 846-1); more recent studies of the Finley quarry exposure suggest Holocene faulting (Sherrod and others, 2016 #7690 Site 3). The 10-m-high exposure of Columbia River Basalt Group (13.5 Ma) is unconformably buried by late Pleistocene flood deposits. The section is interpreted as being complexly faulted; Sherrod and others (2016 #7690) presents evidence of paleoearthquakes in the last 30 k.y., including two earthquakes after the late Pleistocene Missoula floods, based on $^{230}\text{Th}/\text{U}$ series ages on these pedogenic carbonates. The timing of the most recent coseismic surface deformation according to Sherrod and others (2016 #7690) post dates the deposition of Mazama ash; the event occurred prior to Mazama ash deposition. A preserved scarp on basalt in the outcrop suggests 1.8–2.8 m of post-flood vertical displacement. The scarp is interpreted to reflect coseismic deformation based on airborne and ground magnetic data that indicates that the northernmost strand of the Wallula fault zone passes directly through Finley quarry (Sherrod and others, 2016 #7690). GSA Data Repository 2016116 contains a composite log of prior studies of Finley quarry.

Site 846-2. The east wall of a gully exposure of the Umapine fault west of Milton Freewater was examined and logged by Mann and Meyer (their site A, 1993 #35). The gully exposed a several meter wide fault zone consisting of several fault strands dipping 55–63° to the north. Brecciated basalt of the Miocene Columbia River Basalt Group was exposed in the footwall. The hanging wall consisted of a 5-m-thick sequence of ash-bearing gravelly colluvial horizons, interbedded with poorly consolidated sandstones. Samples of volcanic ash from the hanging wall section have geochemical fingerprints

	<p>that are correlative with the Mount St Helens "J" ash, some have fingerprints correlative with Mount St Helens "M", and others are intermediate between these tephra. The Mount St Helens "J" and "M" ashes are thought to have been deposited 10.7 ka and 20.35 ka (Mullineaux, 1986 #3773). The footwall section consists of possibly discrete packages of fault-derived colluvium, separated by deposits of loess at intervals of 1–2 m. Mann and Meyer (1993 #3535) interpret this as evidence of four separate Holocene surface-rupturing events with total vertical displacement of about 5 m.</p>
<p>Geomorphic expression</p>	<p>Most faults in the Wallula fault system have youthful geomorphic expressions; they form topographic and vegetation lineaments, scarp-like lineations, and grabens that offset the youngest geologic units (upper Pleistocene to Holocene?) in the area (Kienle and others, 1979 #3728). Mann and Meyer (1993 #3535) describe numerous youthful fault scarps, lineaments, and laterally offset drainages along the fault zone. McQuarrie (1993 #4337) describes clastic dikes of Pleistocene Missoula Flood sediments (Touchet beds) cutting basalt bedrock and fault breccia and appear to be genetically related to faulting along the Wallula Gap fault. Weldon and others (2001 #5648) observed lineaments across Quaternary deposits on 1:100,000-scale DEM of the area.</p>
<p>Age of faulted surficial deposits</p>	<p>Several faults in the Wallula fault system may offset upper Pleistocene to Holocene surficial deposits (Kienle and others, 1979 #3728; Rigby and Othberg, 1979 #3733; Piety and others, 1990 #3733). One exposure of the Umapine fault offsets loess and colluvium containing the Mount Saint Helens "J" ash system (Mann and Meyer, 1993 #3535) which is thought to have been deposited about 10.7 ka (Mullineaux, 1986 #3773); however, Coppersmith and others (2014 #7402) dispute their conclusions. McQuarrie (1993 #4337) describes clastic dikes of Pleistocene Missoula Flood sediments (Touchet beds) that may be related to faulting. Farooqui and Thoms (1991 #5824) describe evidence for late Pleistocene offset of colluvial deposits along a striking fault at Finley Quarry, near the apparent northwest end of the Wallula fault system.</p>
<p>Historic earthquake</p>	
<p>Most recent prehistoric deformation</p>	<p>latest Quaternary (<15 ka)</p> <p><i>Comments:</i> Age-category assignment based on ORActiveFaults (http://www.oregongeology.org/arcgis/rest/services/Public/ORActiveFaults/MapServer downloaded 06/02/2016). Mann and Meyer (1993 #3535) report one or more events that post-date the 10.7 ka age of the Mount Saint Helens "J" ash on the Umapine fault west of Milton-Freewater. The exposure was re-examined by Coppersmith and others (2014 #7402), and they conclude faulting of Holocene sediments appears to be due to slumping subparallel to the range front rather than tectonic faulting. The bedrock is characterized by subhorizontal moulins at this location, trends to the northwest and</p>

	<p>not parallel to the east-west trending Wallula fault. They conclude there is no evidence of Holocene faulting on the Wallula fault based on these observations as opposed to observations by Kienle and others (1979 #3728) who describe young faulting along the Little Dry Creek fault, where young (<13 ka) and older (Palouse Formation) loess deposits are apparently offset. Kienle and others (1979 #3728), Rigby and Othber (1979 #3738), Piety and others (1990 #3733), and McQuarrie (1993 #4337) describe several locations along splays of the Wallula fault system where Touchet beds (10 ka) are either offset or involved in faulting-related dike injection. Sherrod and others (2016 #7690, fig. 2) show six locations where Holocene coseismic deformation is interpreted and an additional three location where there is evidence of Pleistocene coseismic deformation. Pezzopane (1993 #3544) and subsequent compilations (Geomatrix Consultants Inc., 1995 #3593; Madin and Mabey, 1996 #3575; Weldo and others, 2002 #5648) show most faults in the Wallula fault system as active in the middle and late Quaternary (<700–780 ka) or Quaternary (<1.6–1.8 Ma). The most conservative age assignment is adopted here until more definitive studies are conducted.</p>
<p>Recurrence interval</p>	<p><i>Comments:</i> Sherrod and others (2016 #7690) provide a chronology of surface deformation from three earthquakes since 35 ka and additional evidence of latest Pleistocene liquefaction in the Finlay quarry exposure. In contrast, the fault is poorly expressed considering the amount of reported vertical offset during the three notable earthquakes.</p>
<p>Slip-rate category</p>	<p>Less than 0.2 mm/yr</p> <p><i>Comments:</i> Very little well-documented slip data are available for structures in the Wallula fault system. Hutter (1997 #5650) measured 310 m of vertical offset and a maximum of 1 km of strike-slip displacement in Miocene rocks across the longest fault in the Wallula fault system; Coppersmith and others (2014 #7402) report similar offsets. Meyer and Mann (1993 #3535) calculate a maximum vertical-displacement rate of 0.47 mm/yr on the Umapine fault, based on vertical offset of about 5 m since deposition of the 10.7 ka Mount Saint Helens ash, and Mann and Meyer (1993) suggest a long term rate of about 0.3 mm/yr. In contrast, Geomatrix Consultants, Inc. (1995) used estimated slip rates of 0.04–0.07 mm/yr for a reverse faulting scenario and rates of 0.08–0.2 mm/yr for a strike-slip faulting scenario in their modeling of earthquake hazards along the entire Wallula fault system; the vertical-displacement rate of 0.47 mm/yr reported by Sherrod and others (2016 #7690) does not include horizontal slip. Low slip-rate category is assigned based on the poor geomorphic expression of most faults in this zone.</p>
<p>Date and Compiler(s)</p>	<p>2017 Stephen F. Personius, U.S. Geological Survey Kathleen M. Haller, U.S. Geological Survey</p>

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